

# Detection and replenishment of missing data in earthquake catalogs

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Based on

1. Zhuang, J., T. Wang, K. KiyoSugi (2019) Detection and replenishment of missing data in marked point processes.

*Statistica Sinica. published online.*

2. Zhuang, J., Y. Ogata, T. Wang (2017) Earth, Planet, and Space.

# 报告提纲

- 模型问题
- 数据问题

# 地震震级模型和b值

# Gutenberg-Richter magnitude-frequency relation

## G-R magnitude-frequency relation

$$\log_{10} N(\geq m) = a - bm$$

$$\begin{aligned} \Pr\{M > m | M > m_0\} &\approx \frac{N(\geq m)}{N(\geq m_0)} \\ &= \frac{10^{a-bm}}{10^{a-bm_0}} = 10^{-b(m-m_0)} \end{aligned}$$

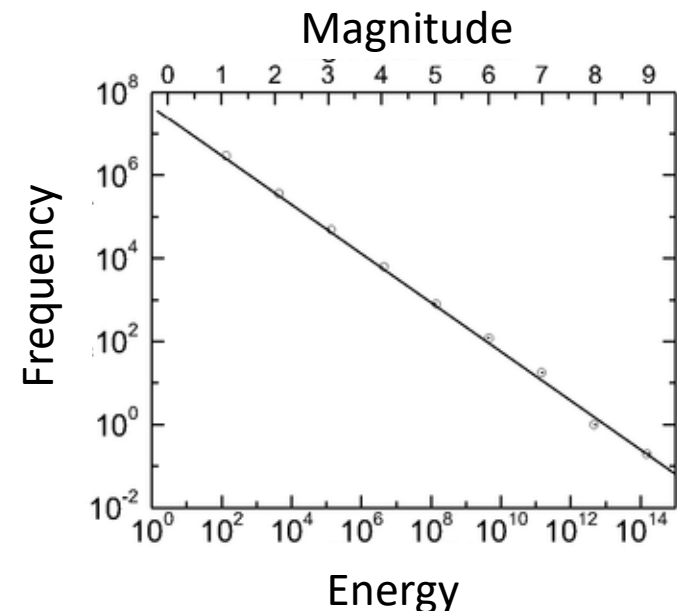
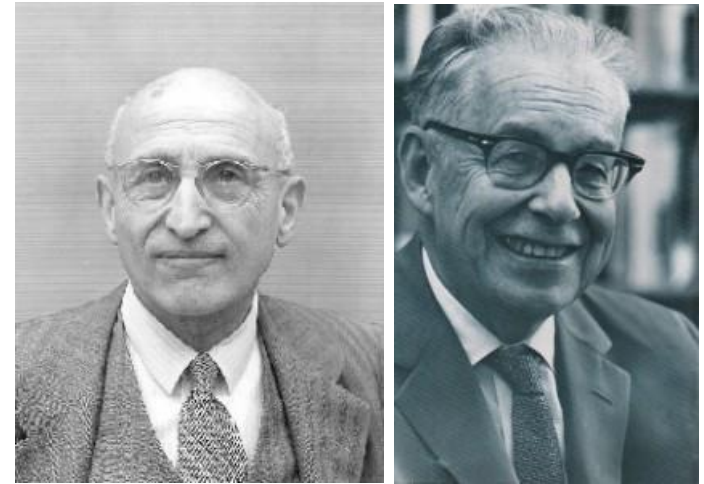
## Probability density function

$$\begin{aligned} f(m) &= b 10^{-b(m-m_0)} \ln 10 \\ &= \beta e^{-\beta(m-m_0)}; \quad m > m_0 \end{aligned}$$

**Power law distribution** for energies, moments and stress drops.

$$\Pr\{E > x\} \sim Cx^{-\alpha}$$

for  $\alpha > 0$ ,  $x > E_0 > 0$ .



## Truncated exponential distribution

$$f(m) = \begin{cases} \frac{\beta e^{-\beta(m-m_0)}}{1 - e^{-\beta(M-m_0)}}, & m_0 \leq m \leq M \\ 0, & \text{otherwise} \end{cases}$$

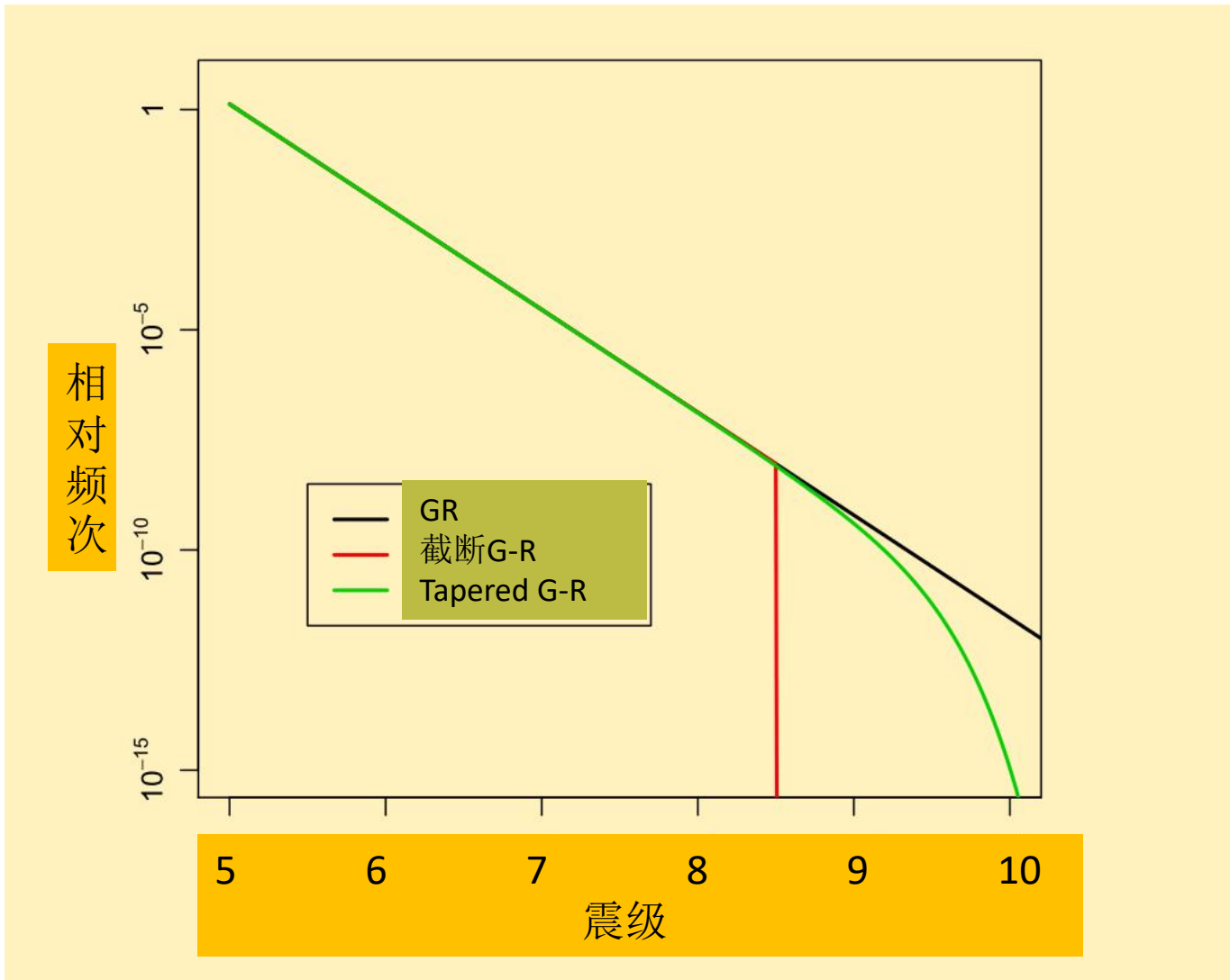
## Tapered exponential distribution

$$f(m) = [\beta + 1.5\gamma 10^{1.5(m-m_c)} \ln 10] \exp[-\beta(m - m_0) - \gamma 10^{1.5(m-m_c)}]$$

$m_c$ : corner magnitude

In moment,  $\Pr\{\text{moment} < x\} = 1 - \left(\frac{x}{S_0}\right)^{-k} \exp\left[\frac{S_0 - s}{S_c}\right]$  **(Kagan) distribution**

$S_c$ : corner moment





# b值的估计和地震震级的模拟



## (1) b-值 最大似然估计

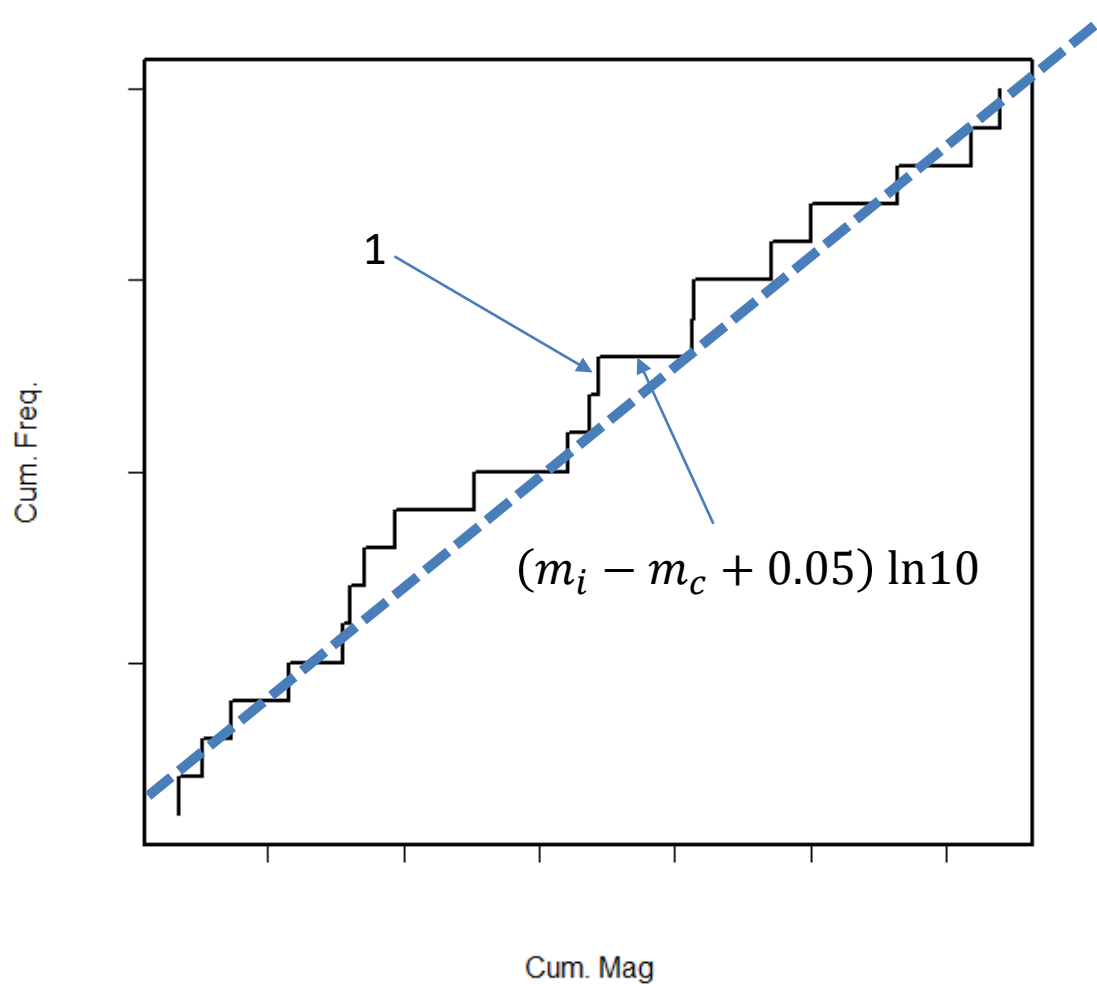
$$\hat{b} = \frac{1}{\left(\bar{M} - m_c + \frac{\Delta}{2}\right) \ln 10}$$

平均震级    震级下限    震级最低位

$$\text{std. err}(\hat{b}) = \frac{\hat{b}}{\sqrt{n}}$$



# (1) b-值 最大似然估计 绘图法

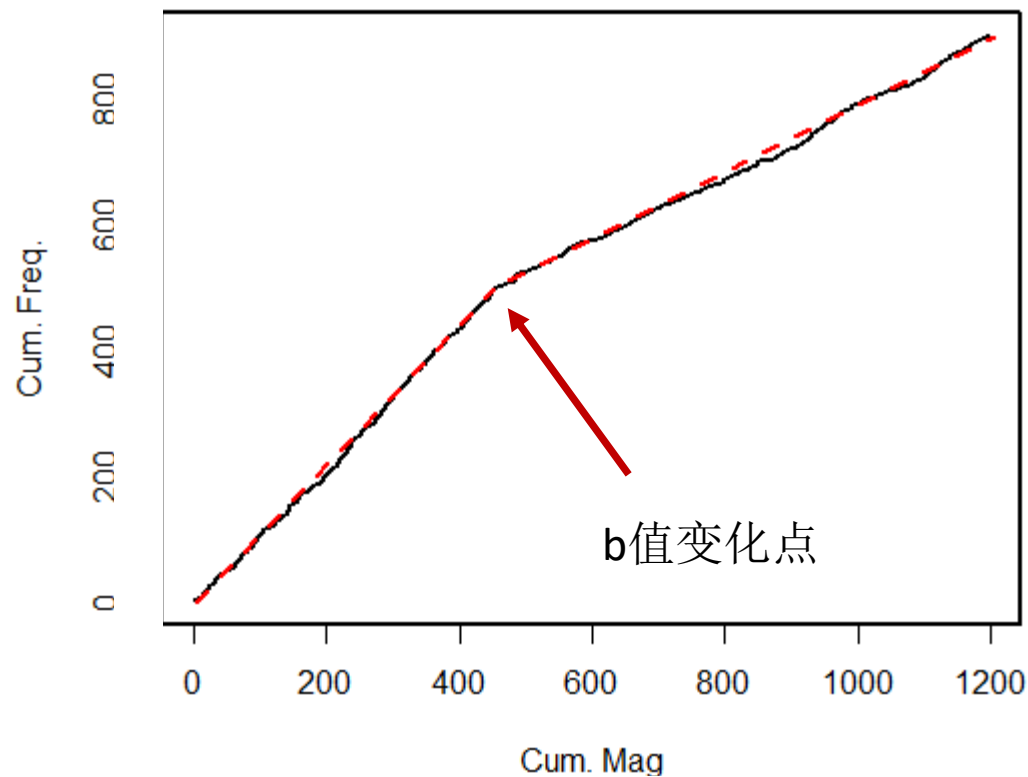


斜率 = b值

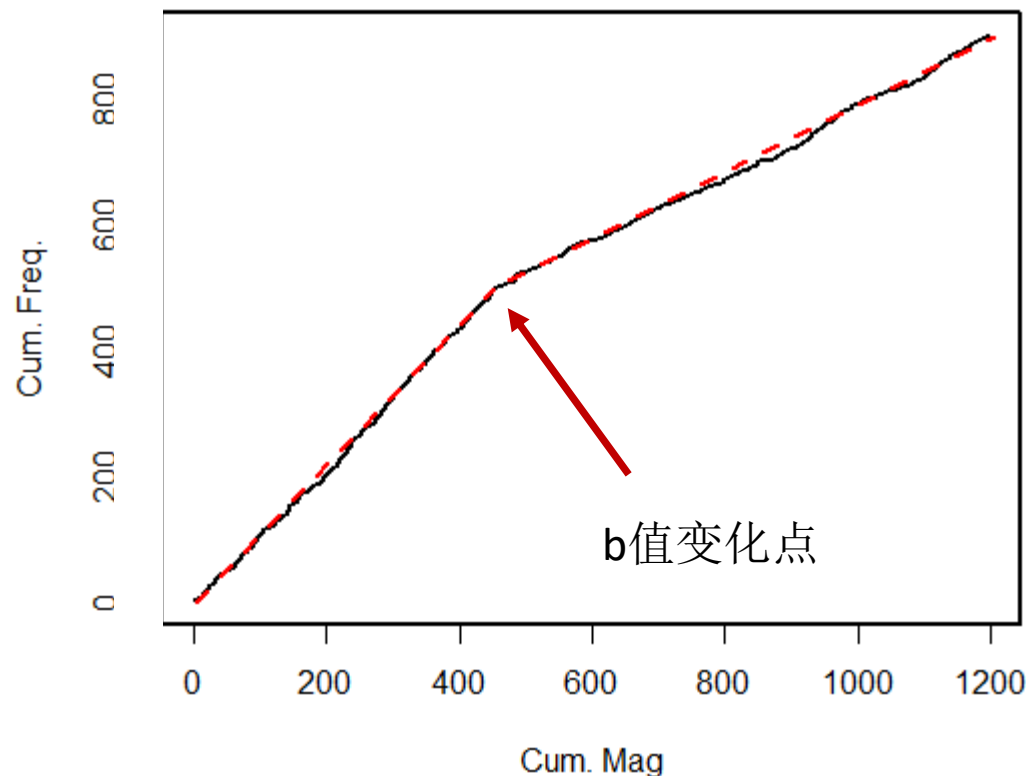
$$i \sim \sum_{k=1}^i m_k$$



## (1) b-值 最大似然估计 绘图法检验b值变化点



## (1) b-值 最大似然估计 绘图法检验b值变化点



把地震分别按照发震时刻、纬度、经度、深度、或者沿断层投影的位置进行排序，就可以看b值在时间、纬度、经度、深度、或者断层上的变化。

# 短期地震活动模型 --- Omori-Utsu公式

## 4. Omori-Utsu formula

- Omori (1894): the rate of felt aftershocks of the 1891  $M_s 8.0$  Nobi earthquake,

$$n(t) = K(t + c)^{-1}, \quad (26)$$

$t$ : the time from the mainshock.

$K$  and  $c$ : constants.

- Utsu (1957): the decay of the aftershock numbers could vary. (Utusu-Omori formula)

$$n(t) = K(t + c)^{-p} \quad (27)$$

$p$ : ranges between 0.6 and 2.5 with a median of 1.1.

- Reasenberg-Jones model

$$\lambda(t, m) = \frac{K s(m)}{(t + c)^p}, \quad (28)$$

$s(m)$ : the magnitude probability density function.

- Multiple Omori-Utsu formula: Not only mainshocks trigger aftershocks, but also large aftershocks may trigger their own aftershocks.

$$\lambda(t) = K/(t - t_0 + c)^{-p} + \sum_{i=1}^{N_T} \frac{K_i H(t - t_i)}{(t - t_i + c_i)^{-p_i}}, \quad (29)$$

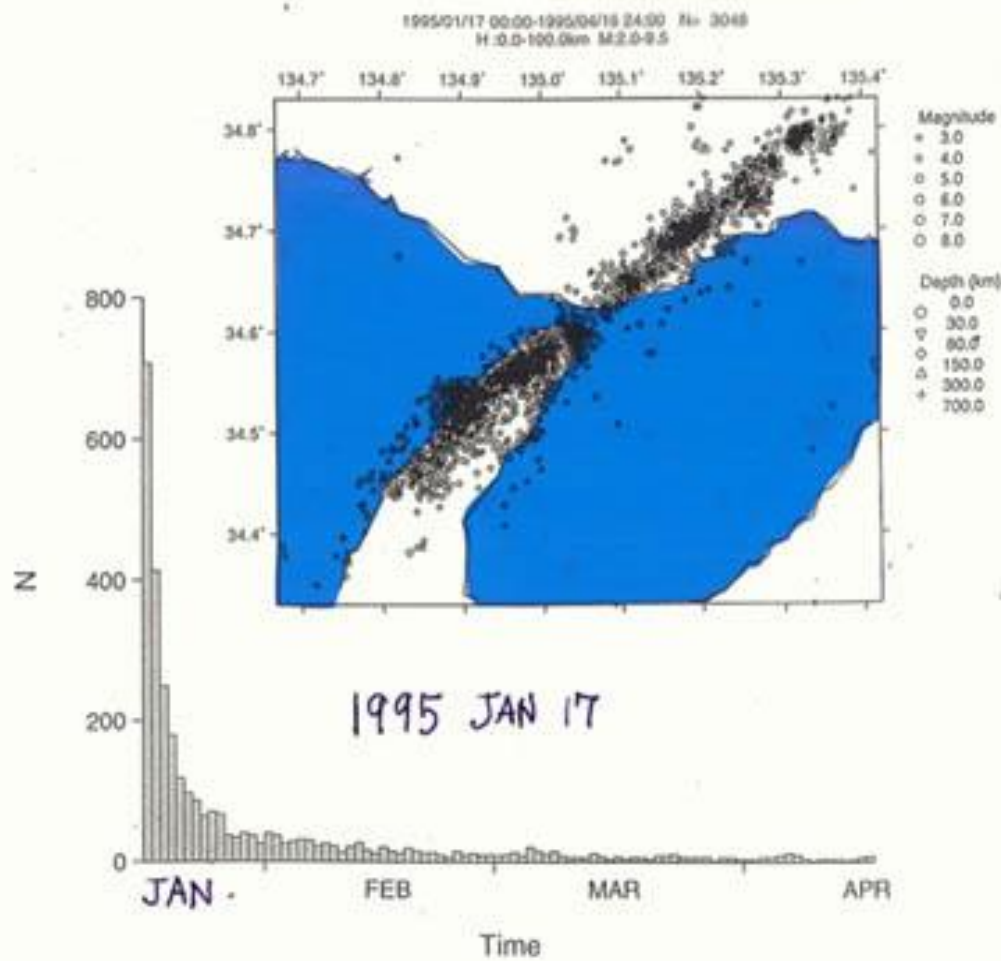
$t_0$ : the occurrence time of the mainshock;

$t_i, i = 1, \dots, N_T$ : the occurrence times of the triggering aftershocks;

$H$ : Heaviside function.

Likelihood function

$$\log L = \sum_{t_i \in [0, T]} \log \lambda(t_i) - \int_0^T \lambda(t) dt$$



$$n(t) = \frac{K}{(t + c)^p}$$

Omori-Utsu Formula (Utsu 1956)

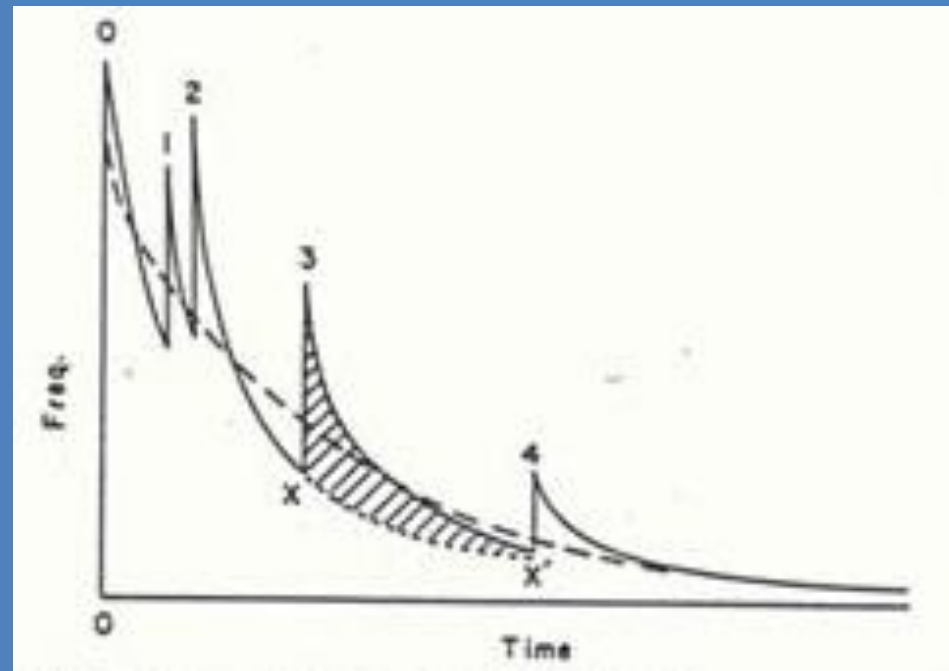
$$n(t) = K(t+c)^{-p} + H(t-T_1)K_1(t-T_1+c_1)^{-p_1} + \dots + H(t-T_n)K_n(t-T_n+c_n)^{-p_n}$$

(Utsu, 1970; Ogata, 1983)

$$\lambda(t) = \mu + \sum_{t_i < t} \frac{K_i}{(t-t_i+c)^p}$$

$$= \mu + K_0 \sum_{t_i < t} \frac{e^{\alpha(M_i - M_0)}}{(t-t_i+c)^p}$$

(Ogata, 1988; 1989)



$\lambda(t)$ : Conditional intensity, hazard function conditioning on the past history



# Temporal ETAS model

## □ Conditional intensity

$$\lambda(t) = \mu + \sum_{i:t_i < t} \kappa(m_i) g(t - t_i)$$

1. Direct productivity:  $\kappa(m) = Ae^{\alpha(m-m_c)}, \quad m \geq m_c$

2. Time p.d.f (Omori-Utsu):  $g(t) = (p-1)(1+t/c)^{-p} / c, \quad t > 0$

## □ Likelihood function

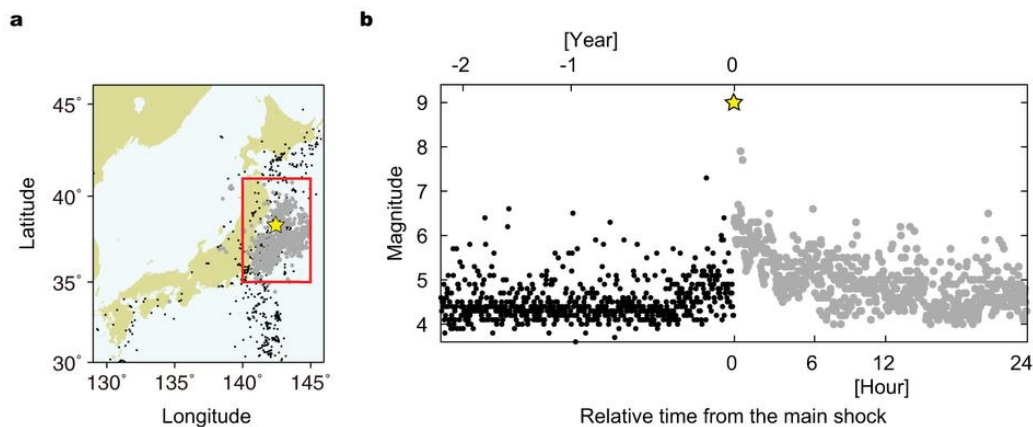
$$\log L = \sum_{t_i \in [0, T]} \log \lambda(t_i) - \int_0^T \lambda(t) dt$$

(Ogata, 1988)

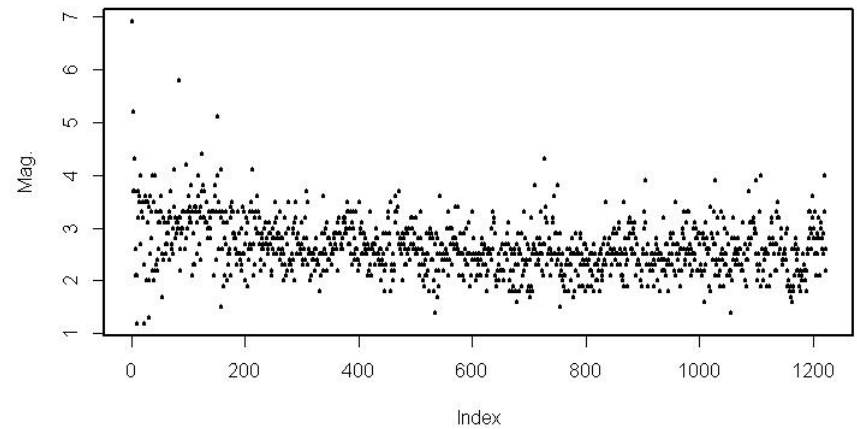
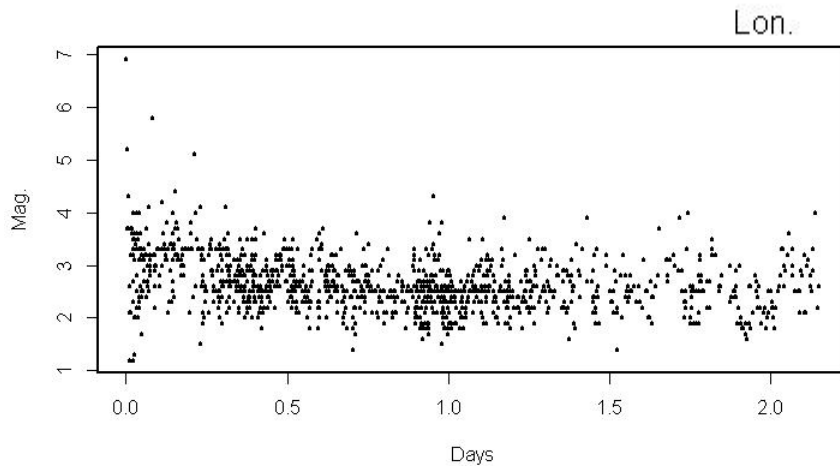
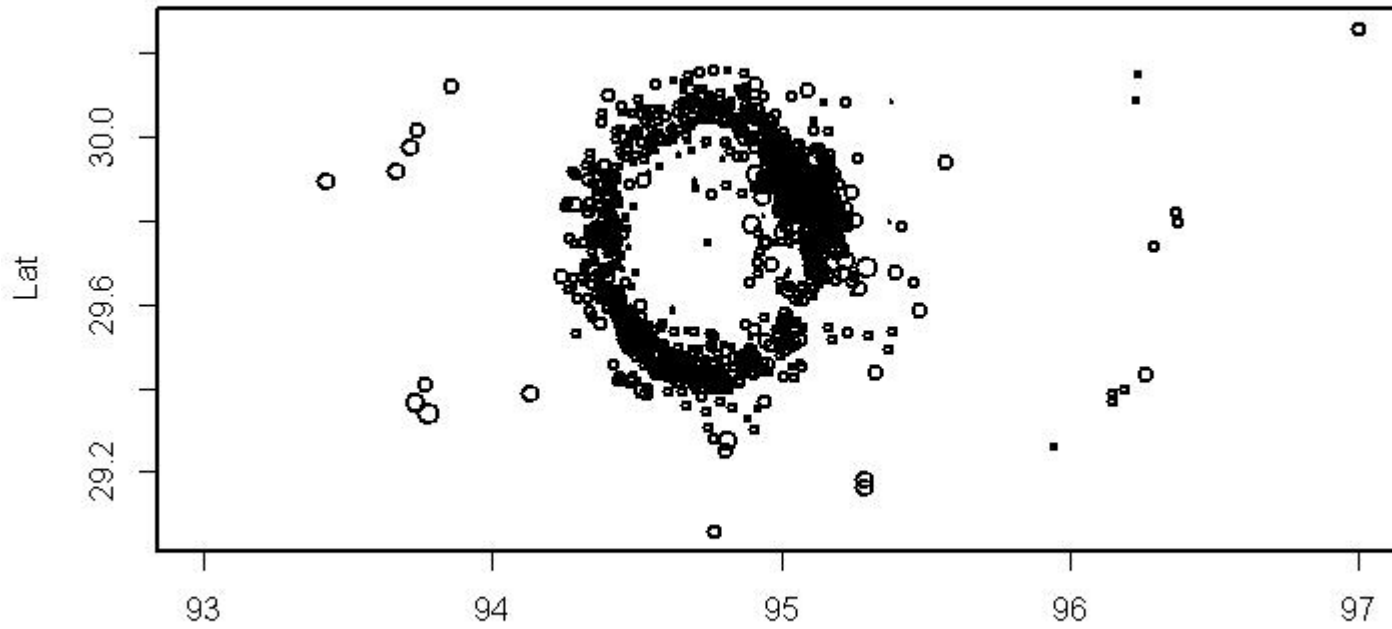
➤ 数据问题      直后余震缺失问题

# Short-term missing of aftershocks

Missing events after the mainshock (from Omi etc, 2013)



# 林芝地震 (2017-11-18 06:34:17.30 29.875 95.057 M6.9)



跳开最震后短时间内的地震进行拟合，结合使用较大震级下限  
 • 汶川地震 M4.0以上

The following is the earthquake probability forecast  
 for the 31th (2008-6-12 14:28 to 2008-6-13 14:28) day after the mainshock

|              | Expected # | Prob.       | Expected.WT | 50%                          | 95%   | 99%    | waiting time |
|--------------|------------|-------------|-------------|------------------------------|-------|--------|--------------|
| M $\geq$ 4.0 | 1.14       | 0.68        | 0.87        | 0.60                         | 2.59  | 3.97   |              |
| M $\geq$ 4.5 | 0.54       | 0.41        | 1.83        | 1.26                         | 5.50  | 8.48   |              |
| M $\geq$ 5.0 | 0.18       | 0.16        | 5.41        | 3.79                         | 16.15 | 24.57  |              |
| M $\geq$ 5.5 | 0.05       | 0.05        | 16.86       | 12.48                        | 46.94 | 72.50  |              |
| M $\geq$ 6.0 | 0.03       | 0.03        | 25.21       | 18.63                        | 71.77 | 117.31 |              |
| 2008-07-15   | 17:26      | 15 Jul 2008 | 09:26       | Mianzhu, Sichuan             | 31.57 |        |              |
|              | 103.98     | 5.0 Ms      |             |                              |       |        |              |
| 2008-07-24   | 03:54      | 23 Jul 2008 | 19:54       | Ningqiang, Shaanxi           | 32.8  |        |              |
|              | 105.5      | 5.6 Ms      |             |                              |       |        |              |
| 2008-07-24   | 15:09      | 24 Jul 2008 | 07:09       | Qingchuan, Sichuan           | 32.82 |        |              |
|              | 105.47     | 6.0 Ms      |             |                              |       |        |              |
| 2008-08-01   | 16:32      | 01 Aug 2008 | 08:32       | Pingwu and Beichuan, Sichuan |       |        |              |
|              | 32.1       | 104.7       | 6.1 Ms      |                              |       |        |              |

# The Kumamoto aftershock sequence data

## Data Selection

Time: 2016/4/1~2016/4/21

Mag.: 1.0+

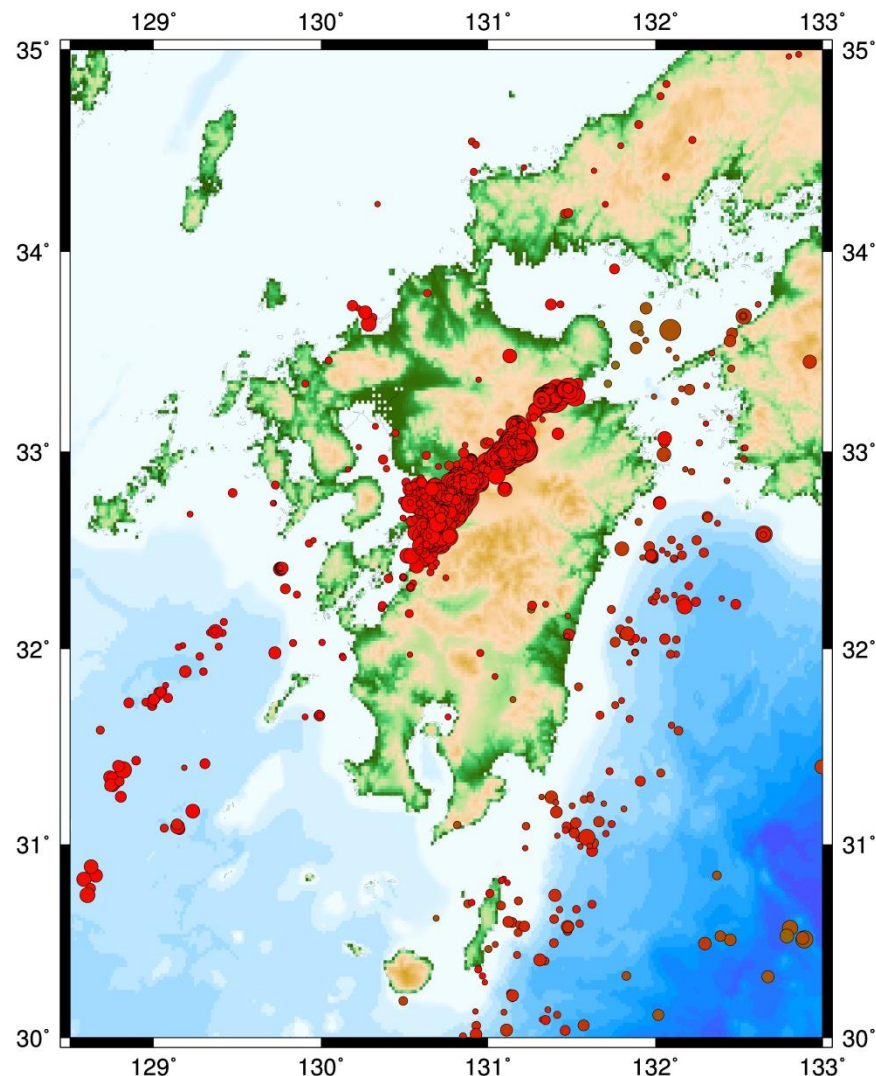
Depth: < 100 km

Space: 128° -- 133°  
30° -- 35°

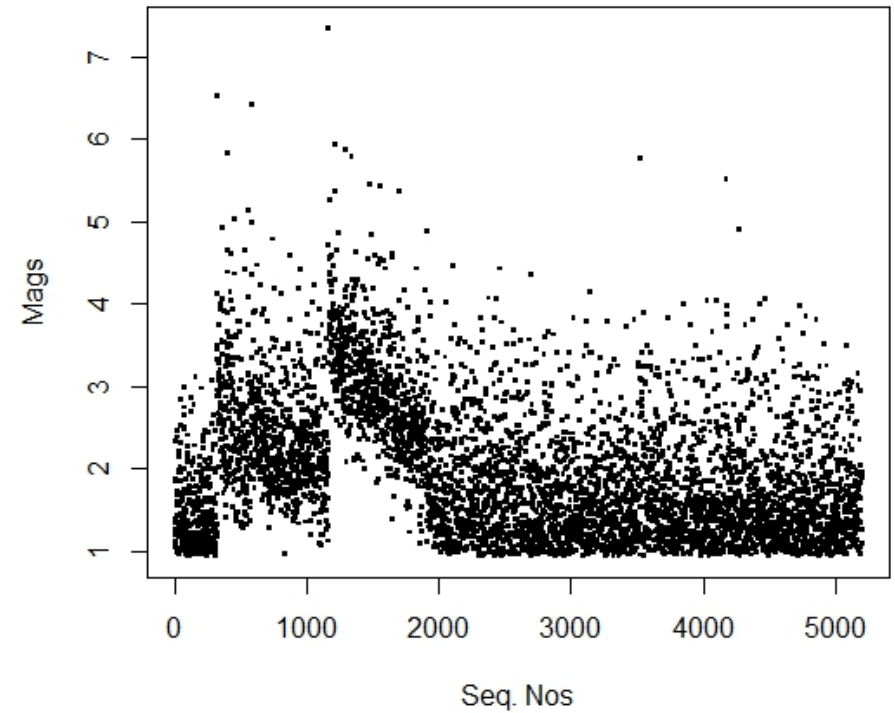
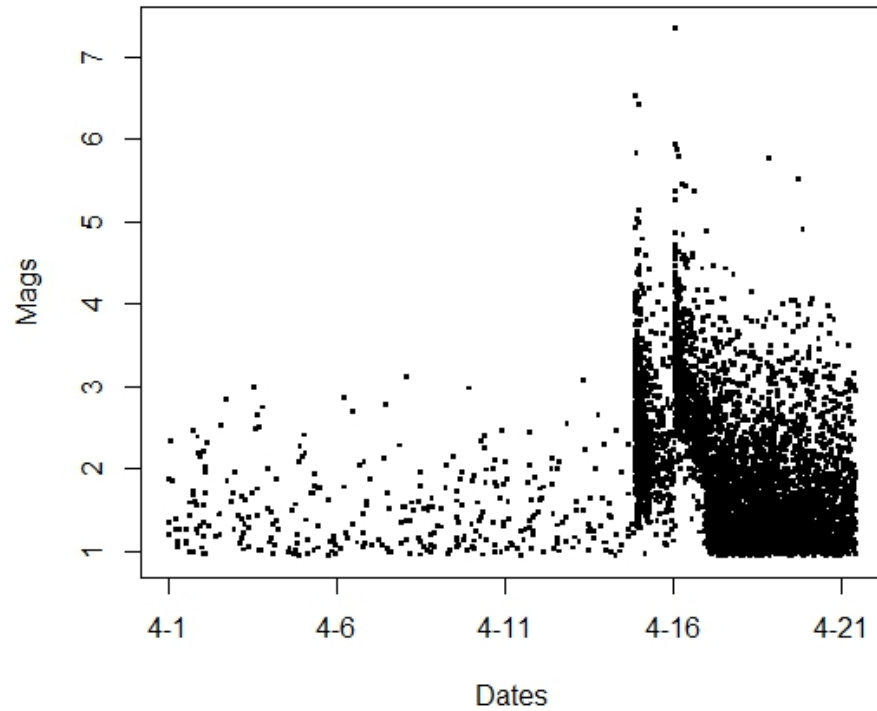
**2016-04-14 21:26 (130.81 32.74) M6.5**

**2016-04-15 00:03 (130.78 32.70) M6.4**

**2016-04-16 01:25 (130.76 32.75) M7.3**



# Application to the recent Kumamoto aftershock sequence data





# Previous studies for fixing the problems of short-term missing aftershocks

## Observational approaches

- ◆ waveform-based earthquake detection methods (e.g., Enescu et al., 2007, 2009; Peng et al., 2007; Marsan and Enescu, 2012; Hainzl, 2016).
- ◆ Energy based description (Sawazaki and Enescu, 2014)

## Statistical approaches

- ◆ (Ogata, Omi, Iwata) Bayesian, assuming GR relation for whole range
- ◆ (Marsan and Enescu, 2012) Assuming Omori-Utsu formula or ETAS model
- ◆ This study: Independence between magnitudes and occurrence times

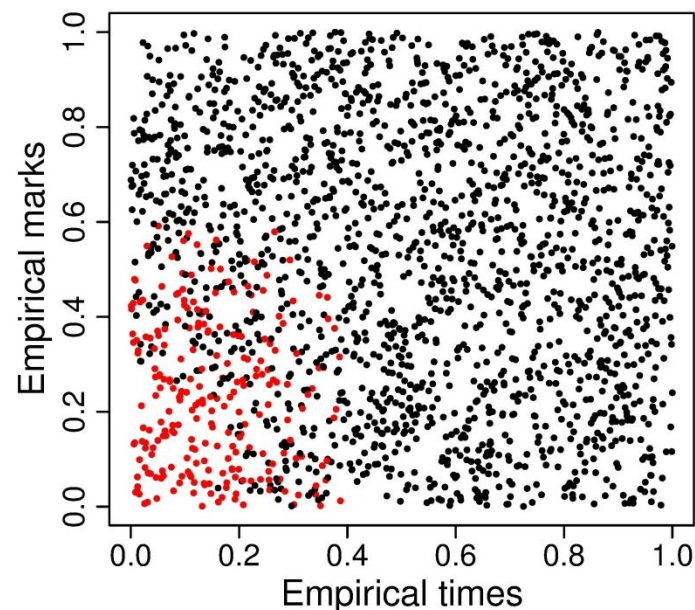
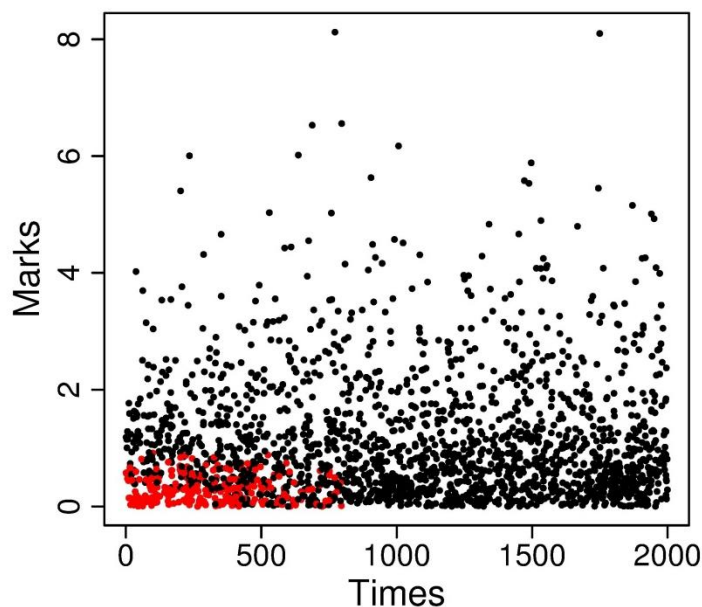
# When data is complete .....

Biscale empirical transformation

$$t_i \rightarrow \tau_i = \frac{i}{N} = \frac{\# \text{ of times } < t_i}{N}$$
$$m_i \rightarrow s_i = \frac{\sum_k \mathbf{1}(m_k < m_i)}{N} = \frac{\# \text{ of marks } < m_i}{N}$$

Original process

→ Homogeneous process

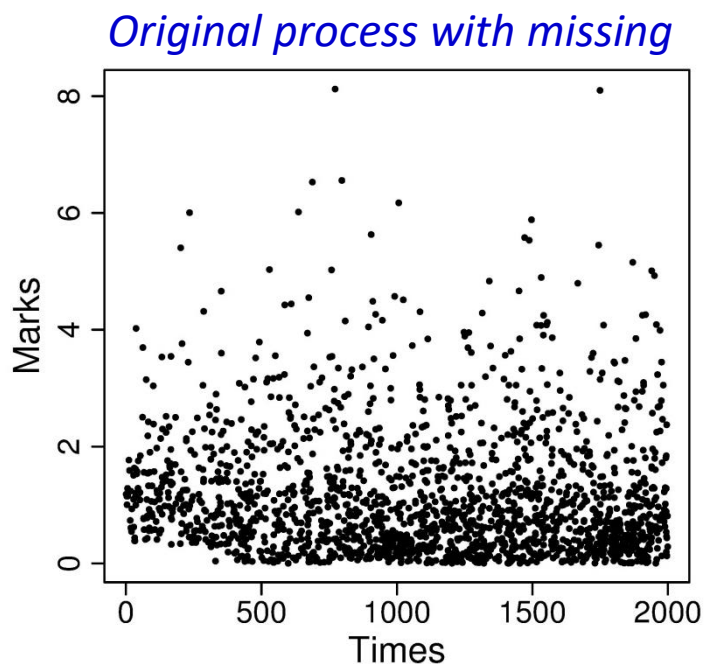


# When data missing exists.....

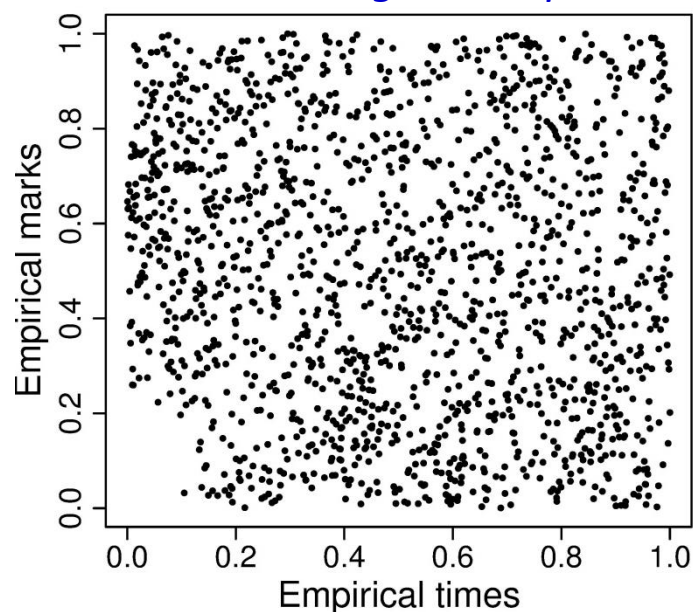
Biscale empirical transformation

$$t_i \rightarrow \tau_i = \frac{i}{N} = \frac{\# \text{ of times} < t_i}{N}$$

$$m_i \rightarrow s_i = \frac{\sum_k \mathbf{1}(m_k < m_i)}{N} = \frac{\# \text{ of marks} < m_i}{N}$$

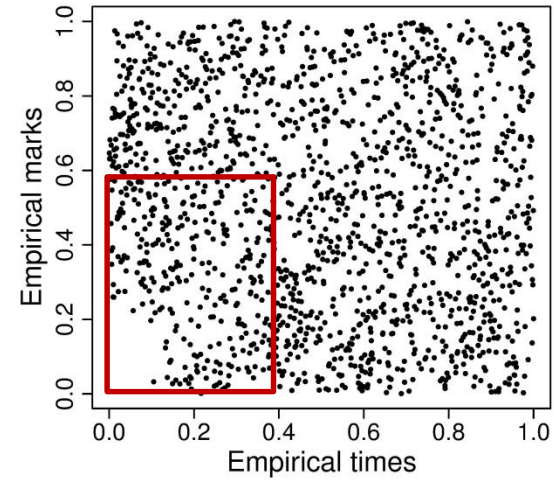
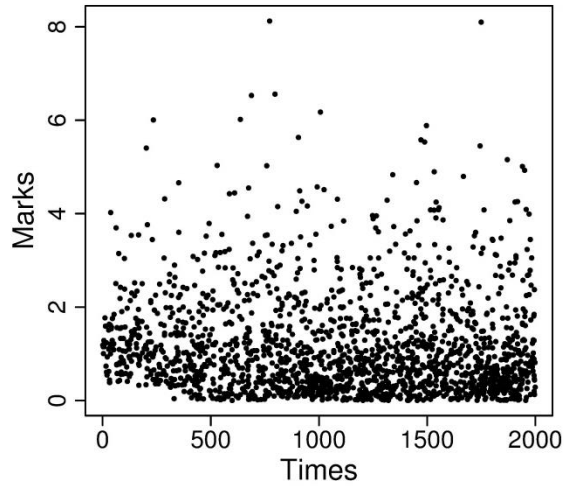


→ *Non-homogeneous process*



# Key points for Replenishing

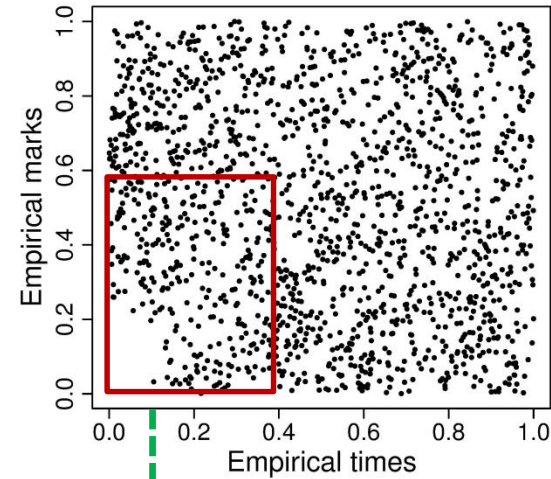
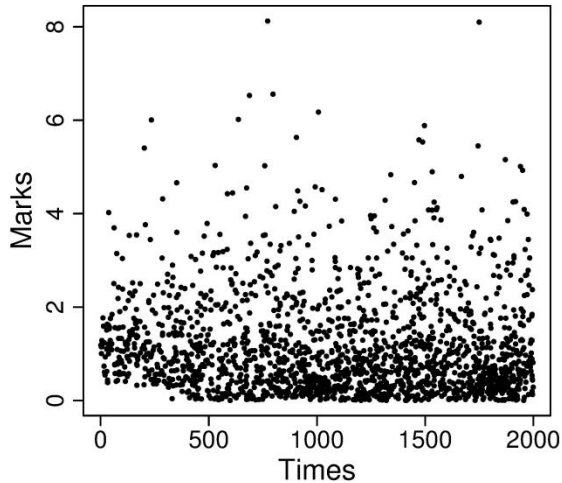
Non-complete dataset



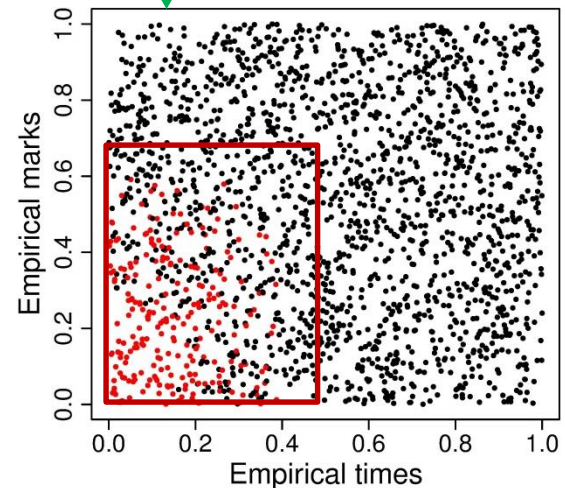
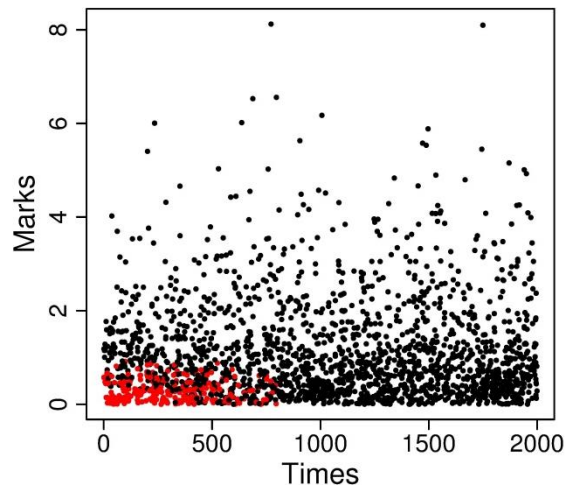


# Key points for Replenishing

Non-complete dataset



Complete dataset

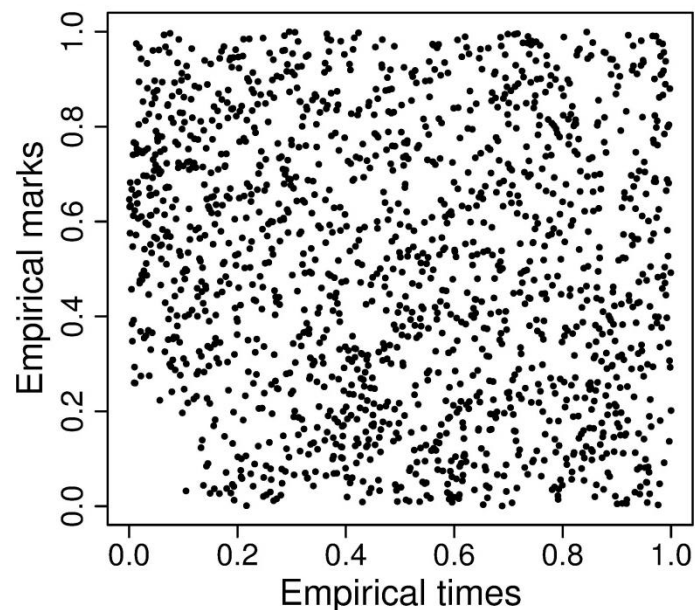
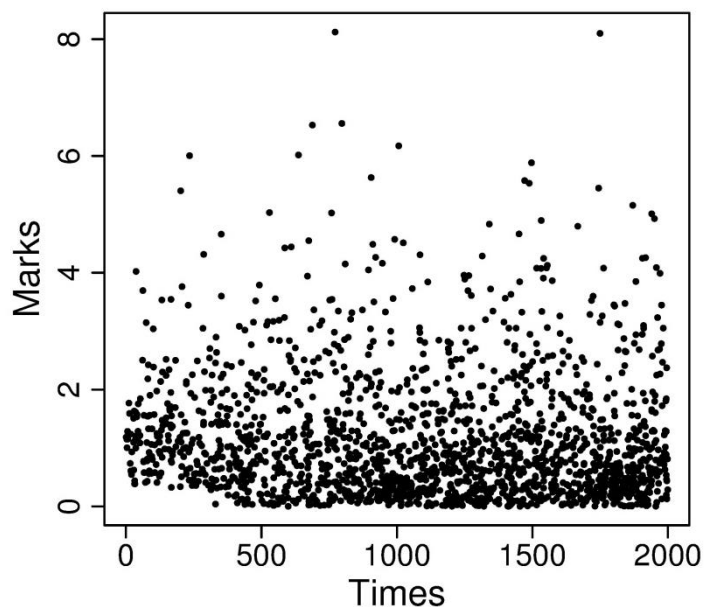


Restore  
*missing area*  
without  
knowing  
missing data  
(red dots)?

# Replenishing algorithm

Step 1. Transform the process using the biscale empirical transformation

$$t_i \rightarrow \tau_i = \frac{i}{N} = \frac{\# \text{ of times } < t_i}{N}$$
$$m_i \rightarrow s_i = \frac{\sum_k \mathbf{1}(m_k < m_i)}{N} = \frac{\# \text{ of marks } < m_i}{N}$$

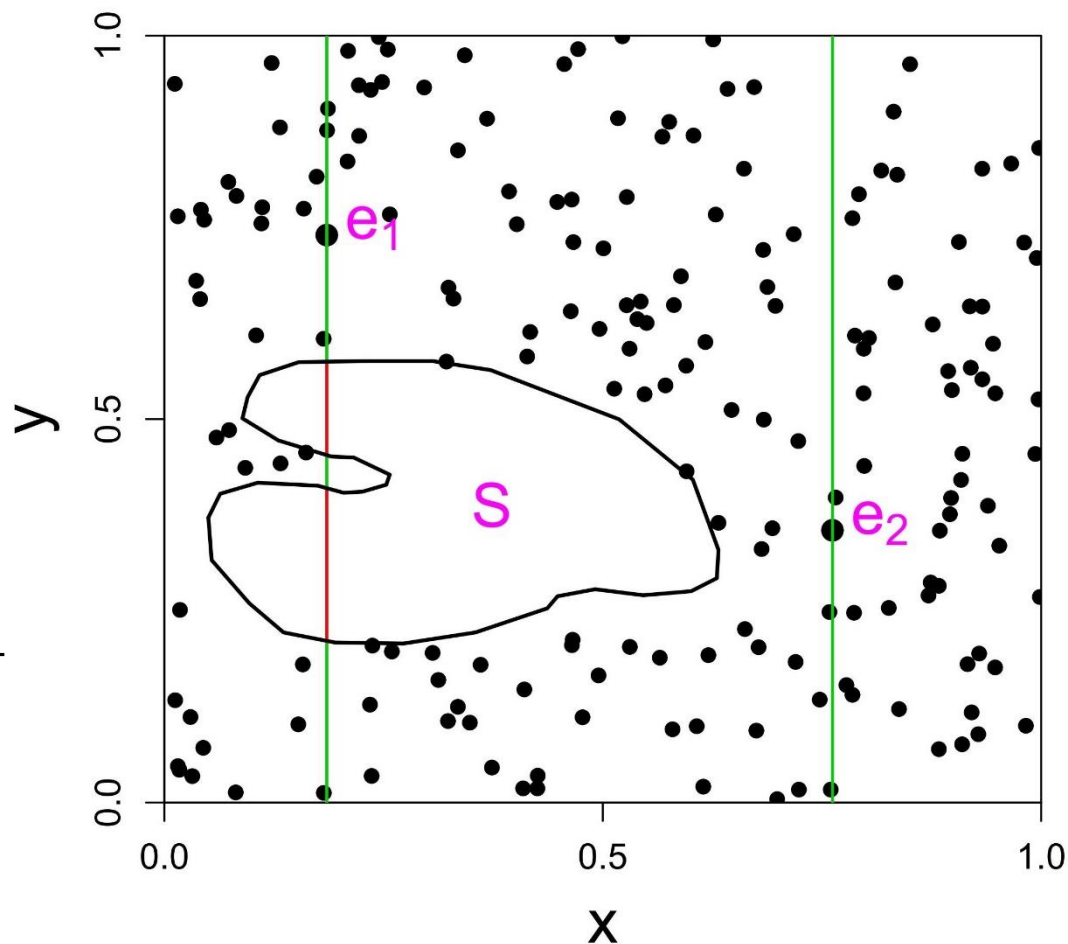


## Heuristic illustration

Heuristic illustration for estimating empirical probability distribution function when missing happens in homogeneous process

$$\widehat{F}_X(x) = \frac{\sum_i w(e_i) I(x_i < x)}{\sum_i w(e_i)}$$

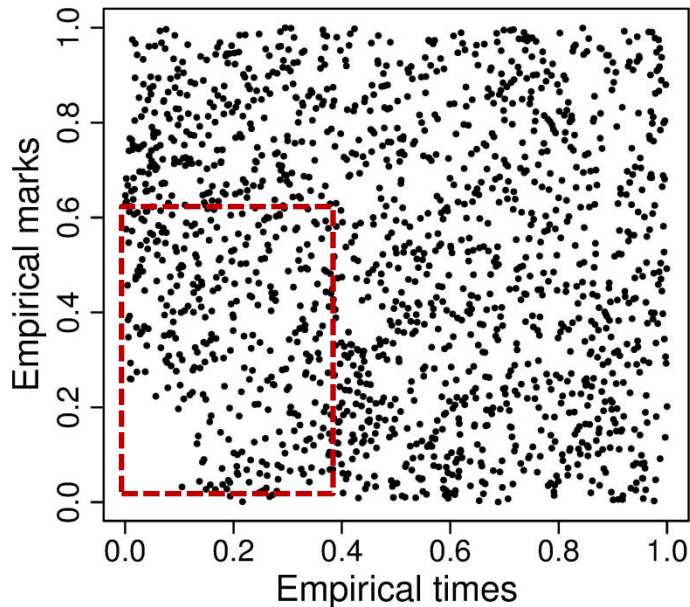
$$w(e_i) = \frac{1}{1 - \int_0^1 I((x_i, y) \in S) dy}$$





# Replenishing algorithm

Step 2. Specify area  $S$  that contains the missing data



The missing area  $S$  satisfies

$$\int_M \mathbf{1}((t, m) \notin S) dF_2(m) > 0$$

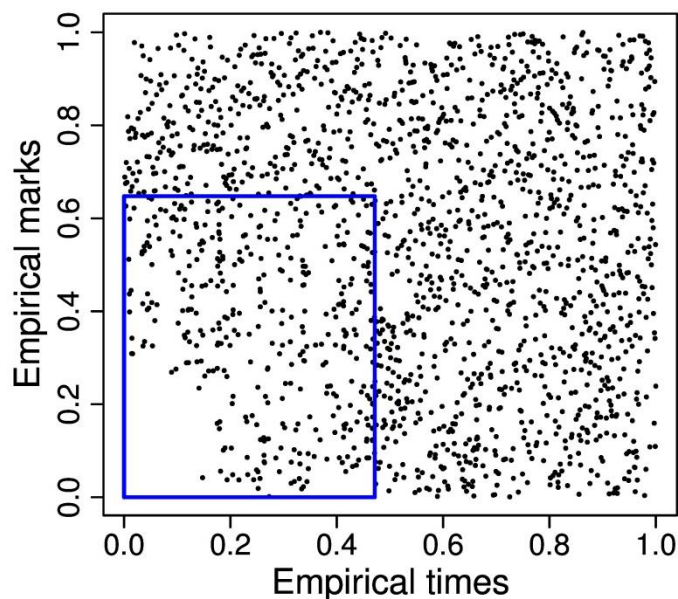
for all  $t \in [0, T]$  and

$$\int_0^T \mathbf{1}((t, m) \notin S) \mu_g(t) dt > 0$$

for all  $m \in M$ .

# Replenishing algorithm

Step 3. Calculate the missing area in the biscale transformation domain based on complete data



$$F_1^*(t) = \frac{\sum_{j=1}^n w_1(t_j, m_j, S) \mathbf{1}(t_j < t)}{\sum_{j=1}^n w_1(t_j, m_j, S)}$$

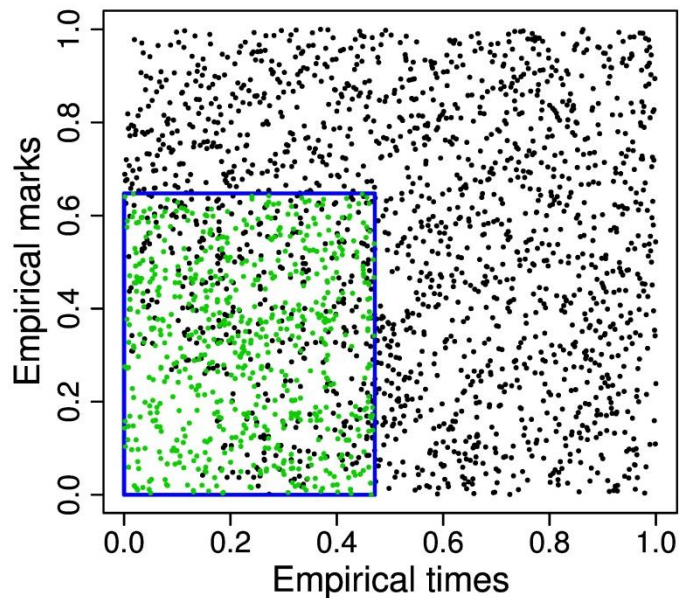
$$F_2^*(m) = \frac{\sum_{j=1}^n w_2(t_j, m_j, S) \mathbf{1}(m_j < m)}{\sum_{j=1}^n w_2(t_j, m_j, S)}$$

$$w_1(t, m, S) = \frac{\mathbf{1}((t, m) \notin S)}{\int_M \mathbf{1}((t, s) \notin S) dF_2^*(s)}$$

$$w_2(t, m, S) = \frac{\mathbf{1}((t, m) \notin S)}{\int_M \mathbf{1}((\tau, m) \notin S) dF_1^*(\tau)}$$

# Replenishing algorithm

Step 4. Generate data point in the missing area



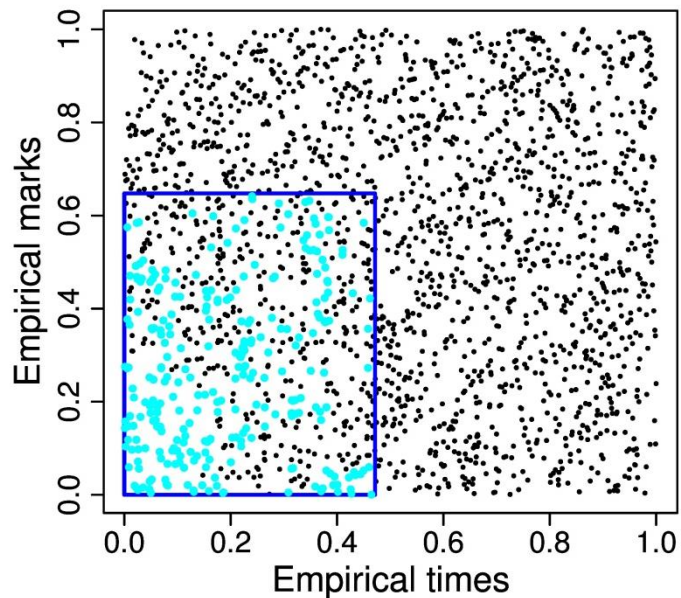
Generate events uniformly distributed in the missing region  $S^*$  (image of  $S$ )

$$\# \text{events} \sim NB(k, 1 - |S^*|)$$

$k$ : # of observed events outside of  $S^*$

# Replenishing algorithm

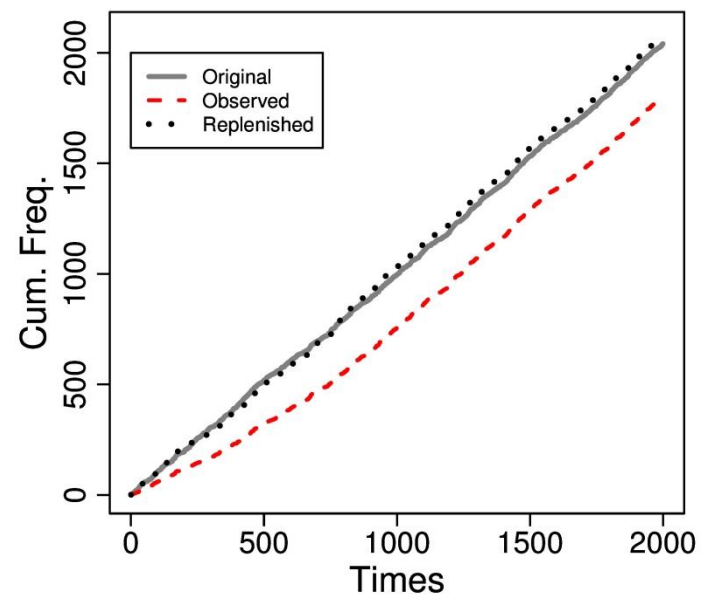
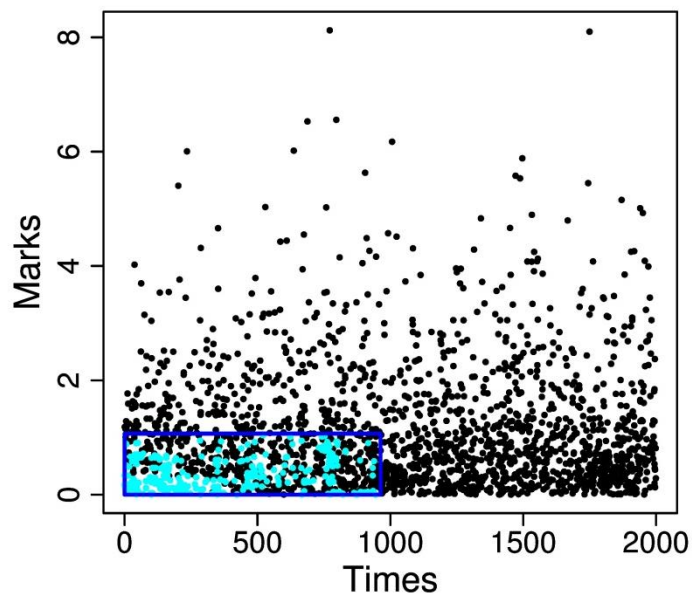
Step 4. Remove sequentially a simulated data point for each existing point in the missing area





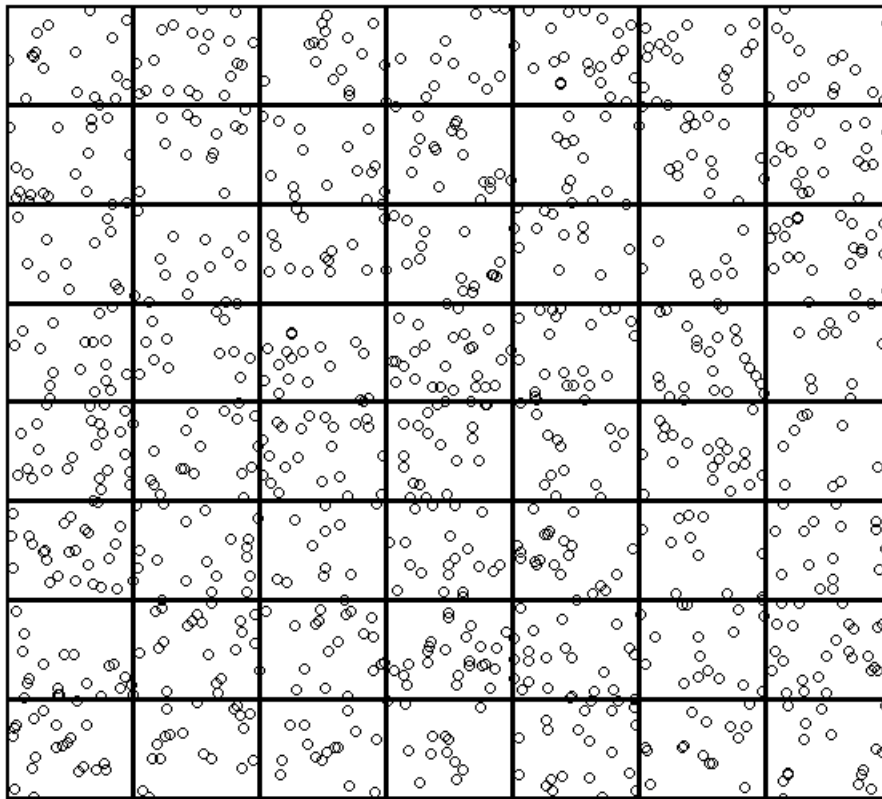
# Replenishing algorithm

Step 5. Transform back all the events into the original domain



# How to test existence of missing: Testing method

$L_2$  columns



$L_1$   
rows

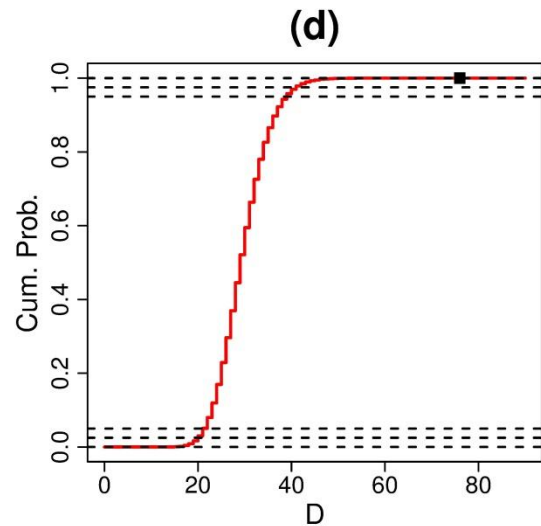
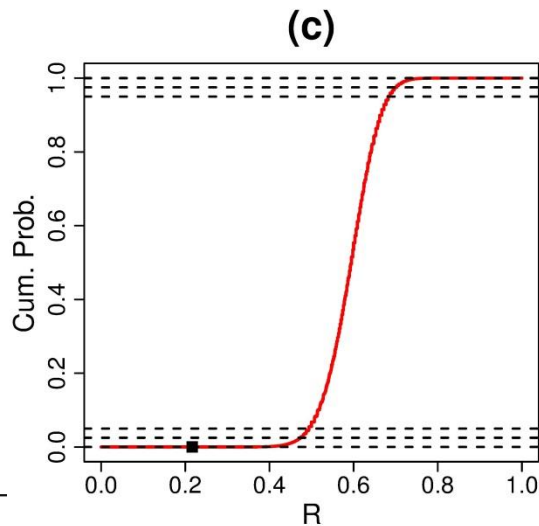
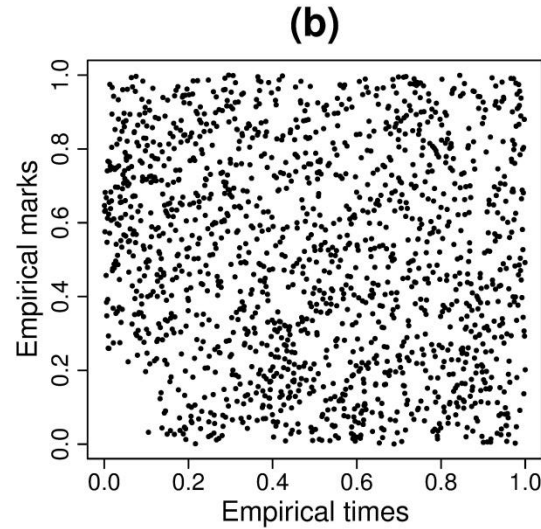
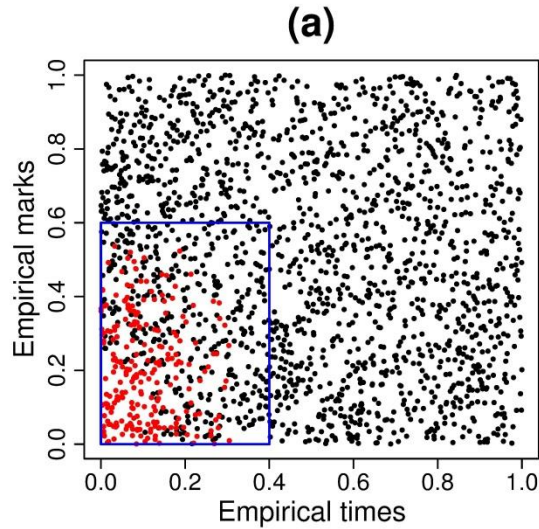
$$R = \frac{\min\{C_1, C_2, \dots, C_L\}}{\max\{C_1, C_2, \dots, C_L\}}$$

$$D = \max\{C_1, C_2, \dots, C_L\} - \min\{C_1, C_2, \dots, C_L\}$$

$C_i$ : #events in cell  $i$

$$L = L_1 \times L_2$$

# Testing method



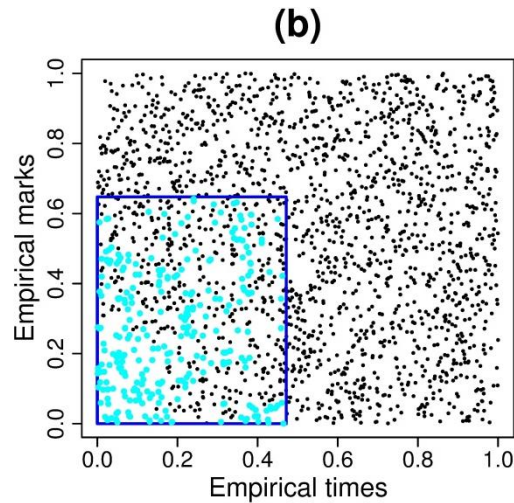
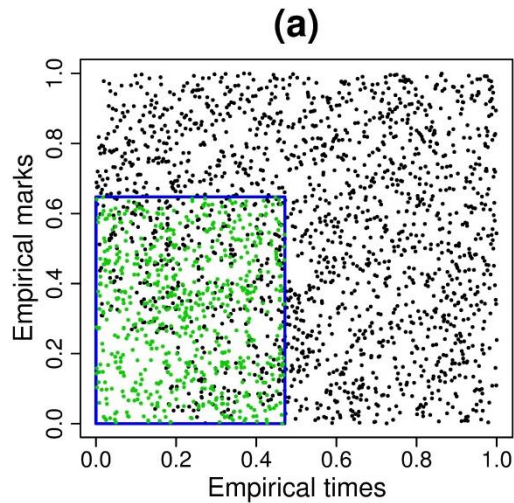
Observed data

$$L_1 = L_2 = 5$$
$$L = 25$$

Distributions of  $R$  and  $D$  when the same number of data points are completely observed

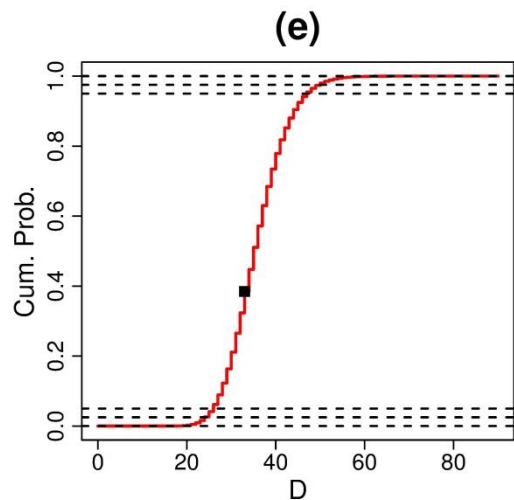
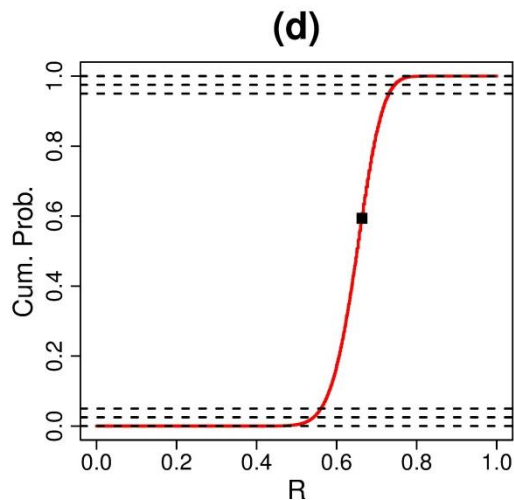


# Testing method



Replenished data

$$L_1 = L_2 = 5$$
$$L = 25$$



Distributions of  $R$  and  $D$  when the same number of data points are completely observed



# More Simulations

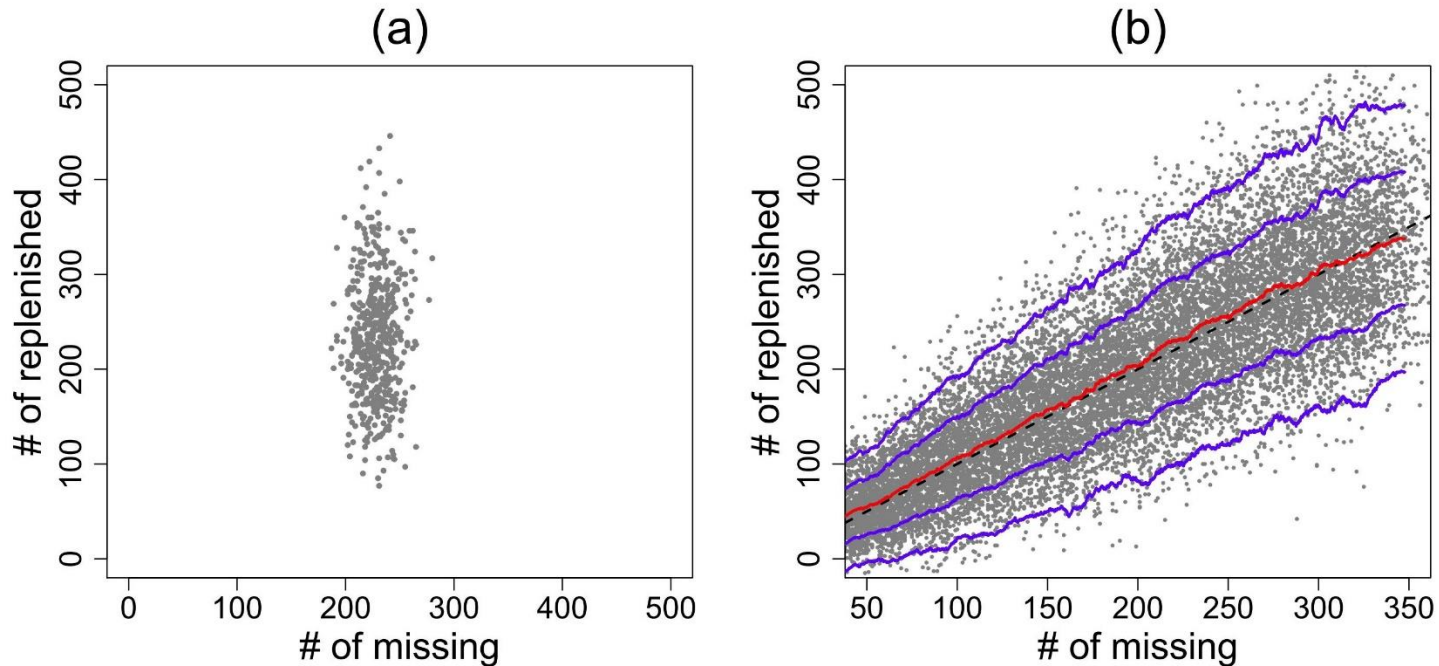
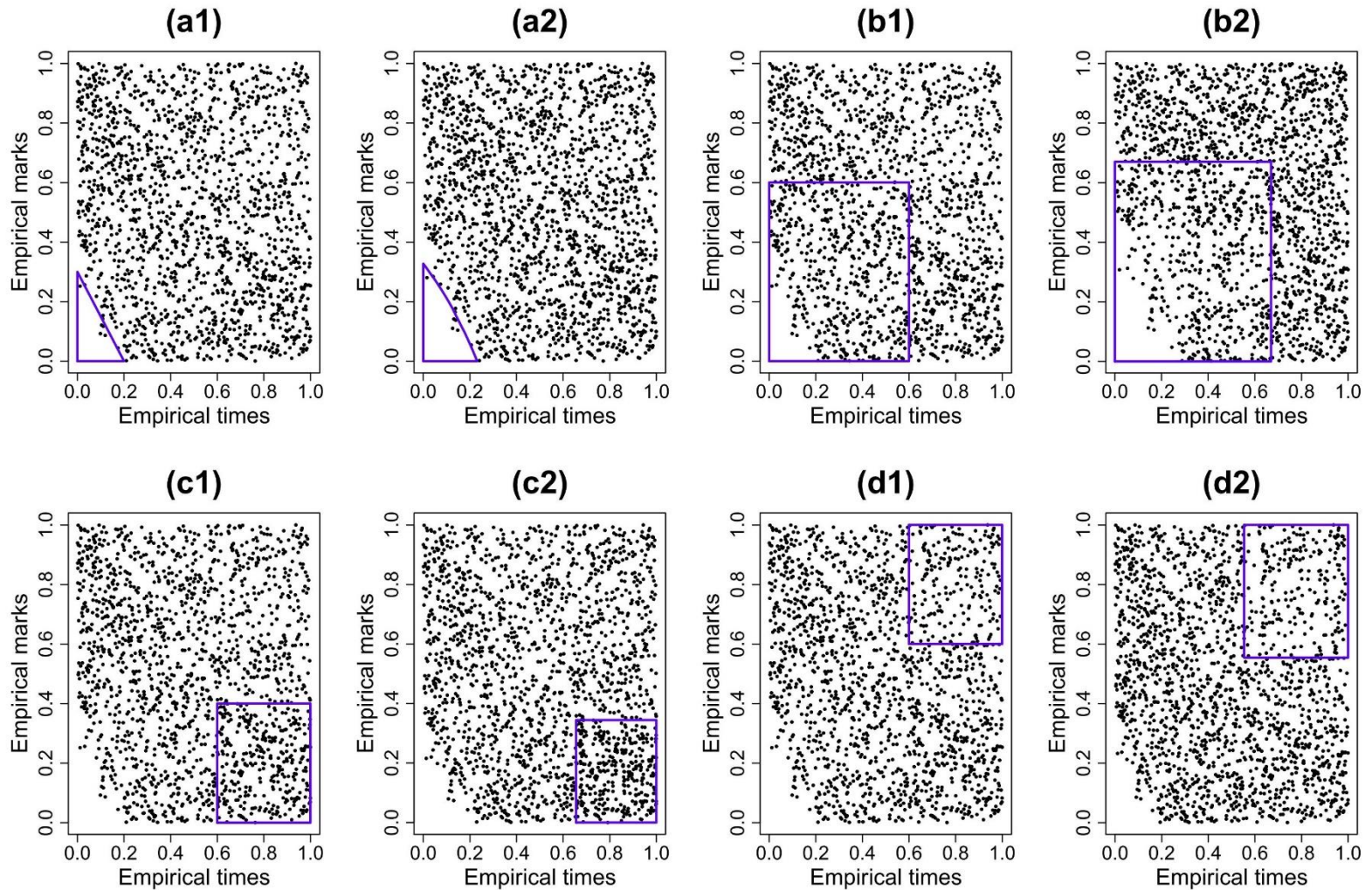


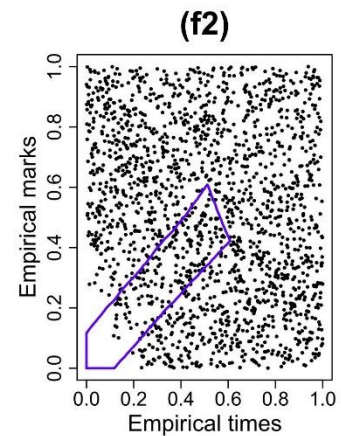
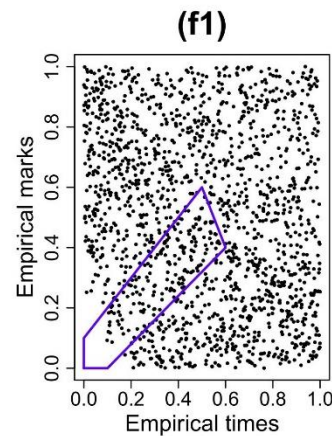
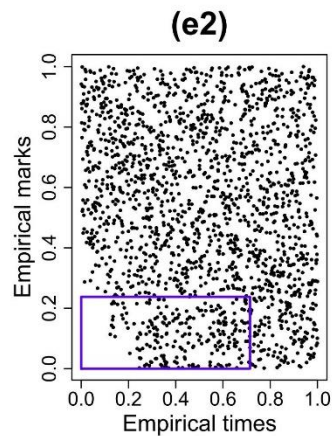
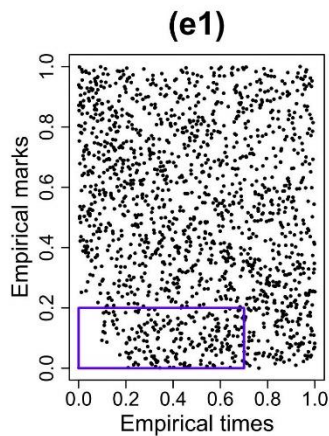
Figure 5: Comparison between the number of true missing events and the number of replenished events. (a)  $\lambda = 2,000$  fixed. (b)  $\lambda$  is drawn from a uniform distribution between 100 and 3000. The dashed lines represent the case where the numbers of missing events and replenished events are equal. The blue and red curves represent the running mean and the corresponding single and double standard deviation bands.



# If wrong selection of missing area (1)



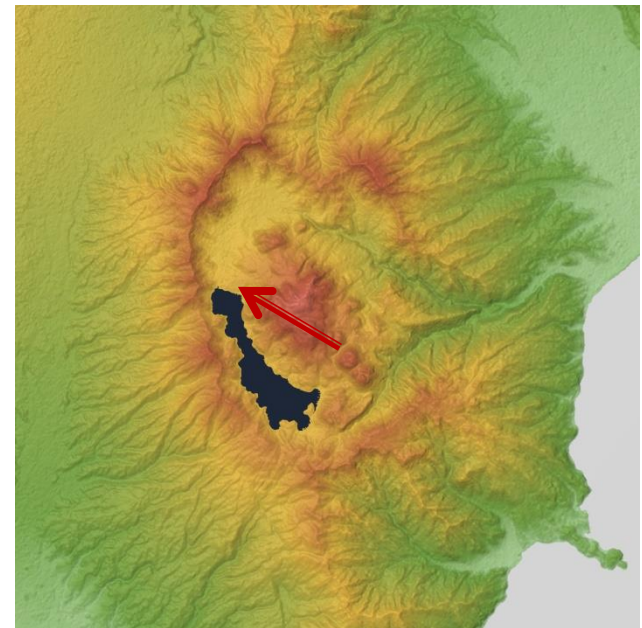
# If wrong selection of missing area (1)

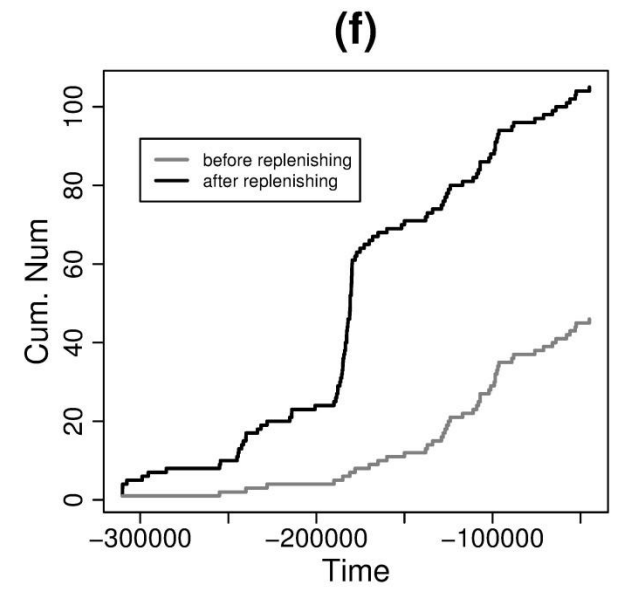
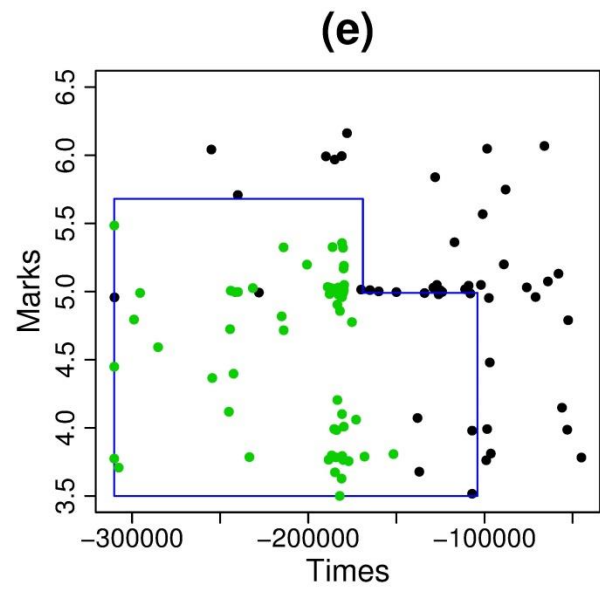
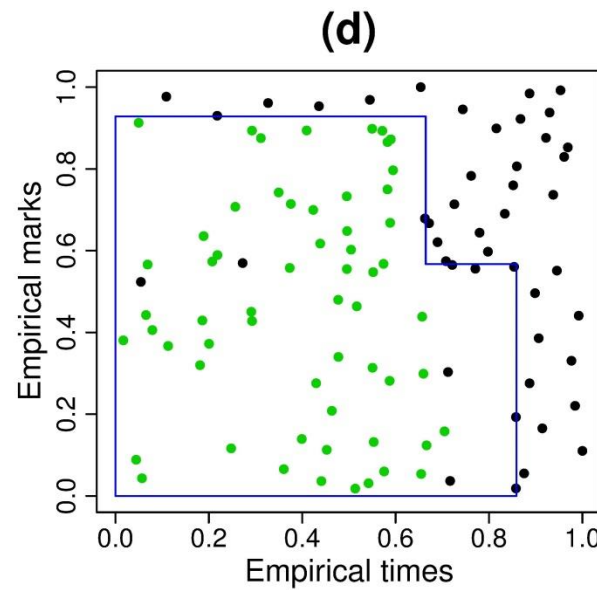
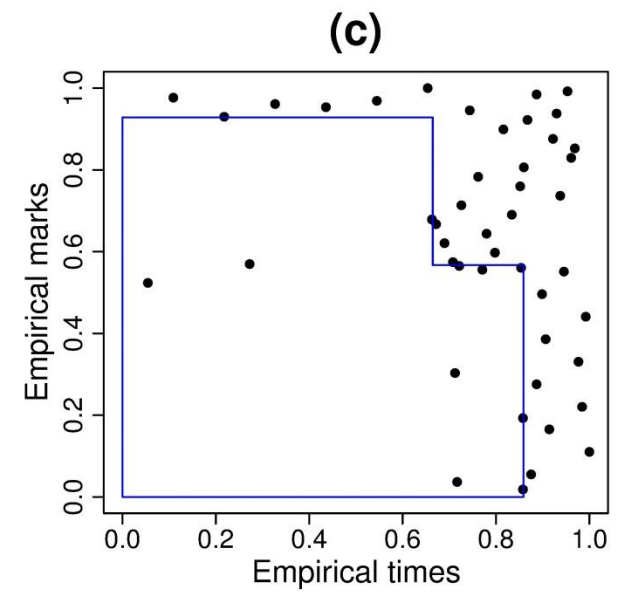
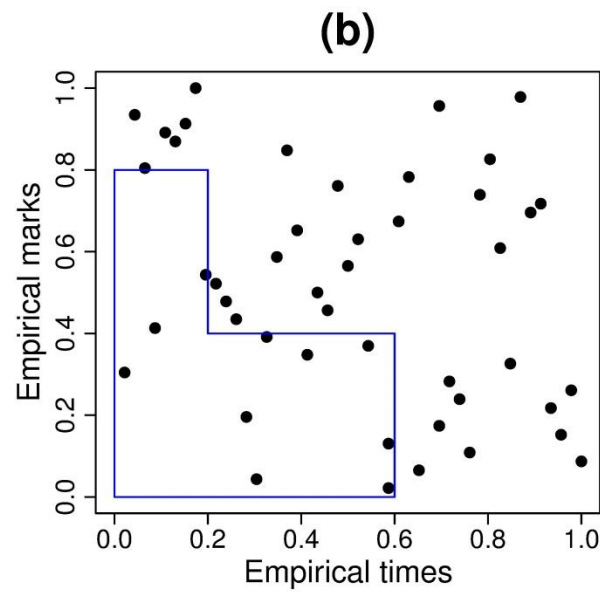
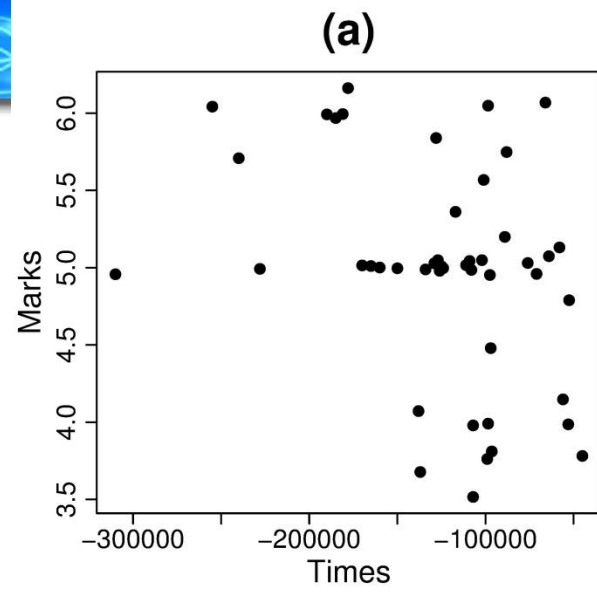




# Application 1: Volcanic eruption record

- Data: eruptions at Hakone volcano
- Data source:
  - Smithsonian's Global Volcanism Program database
  - Large Magnitude Explosive Volcanic Eruptions database (LaMEVE database)
  - additional Japanese databases





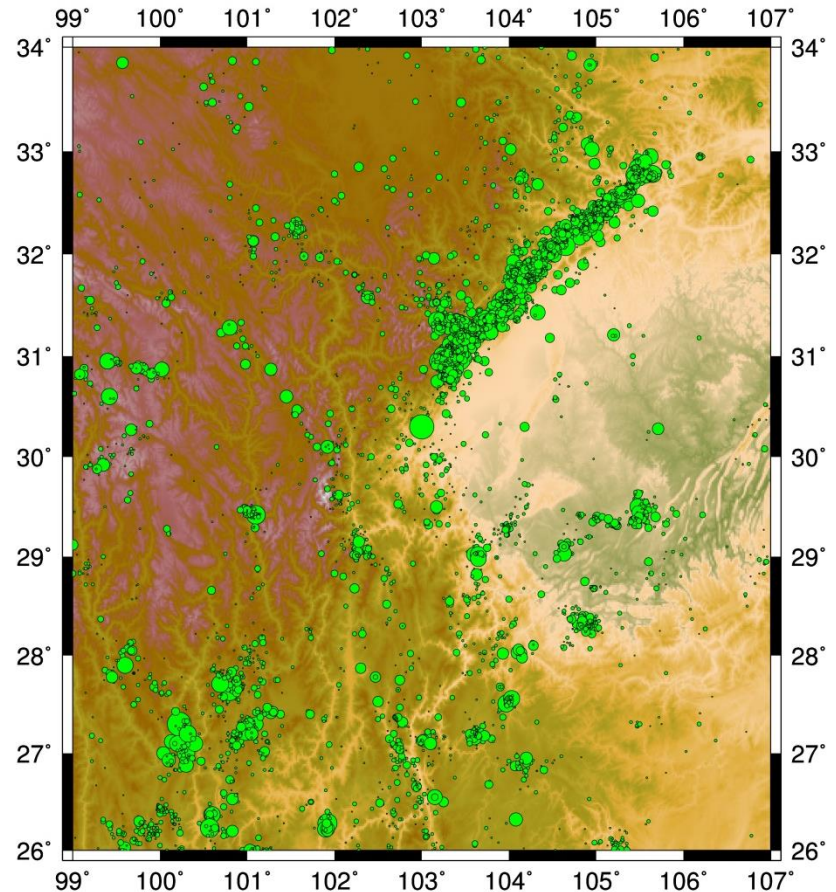
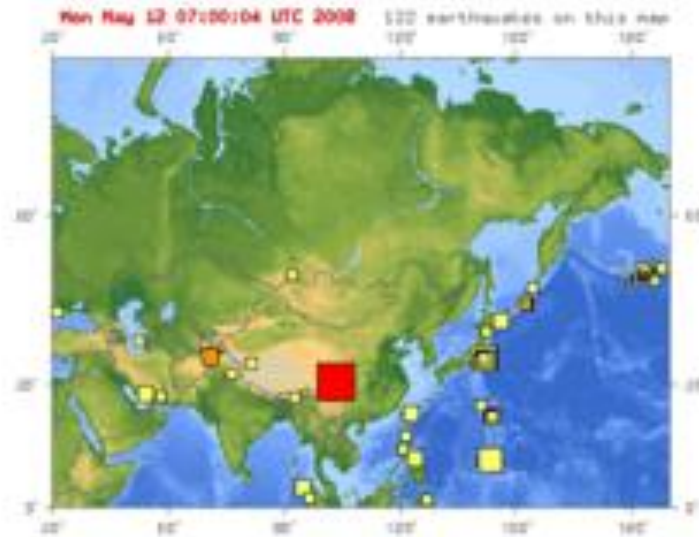
# Application to the Wenchuan aftershock sequence data

## Wenchuan EQ: Mw7.9 (Ms8. 2) 2008/5/12

Data selection:

Time: 1990/1/1~2013/4/20

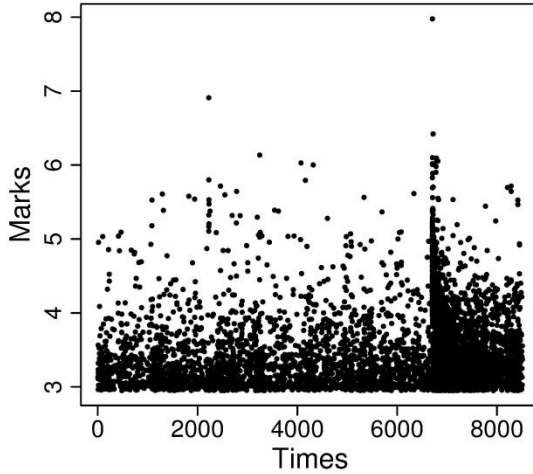
Magnitude: 3.0+



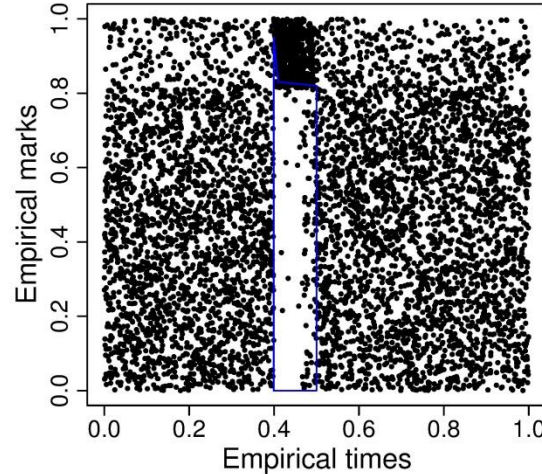


# Application to the Wenchuan aftershock sequence data

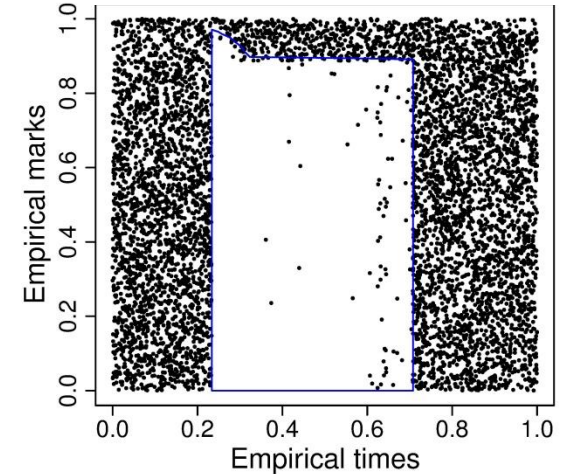
Observed data



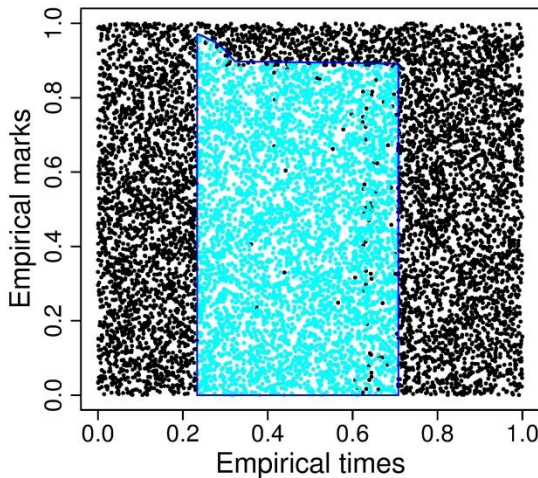
Biscale transformed



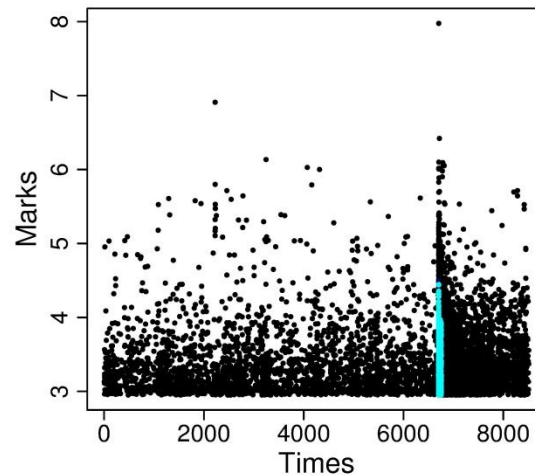
Estimate transformation under complete data



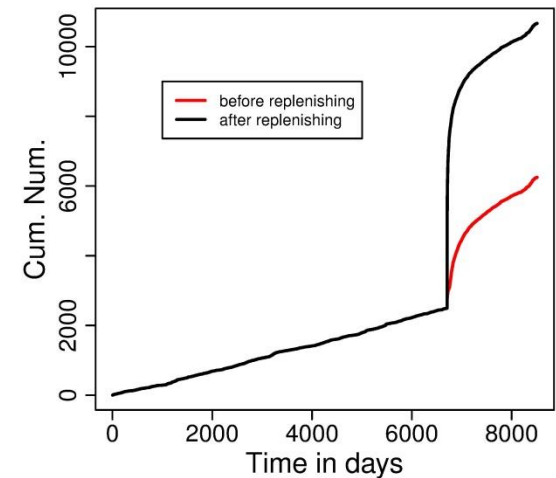
Transformed replenished data



Original replenished data



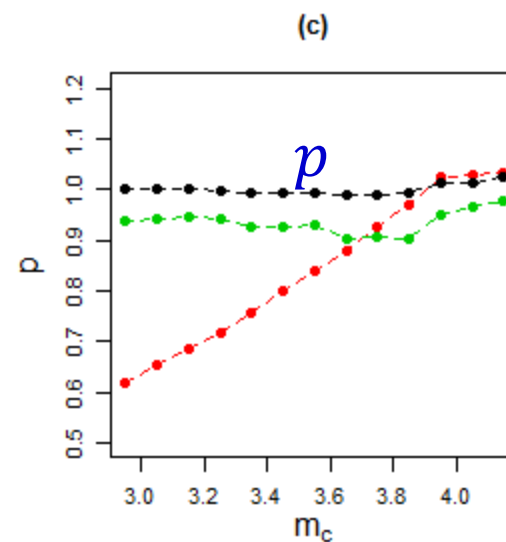
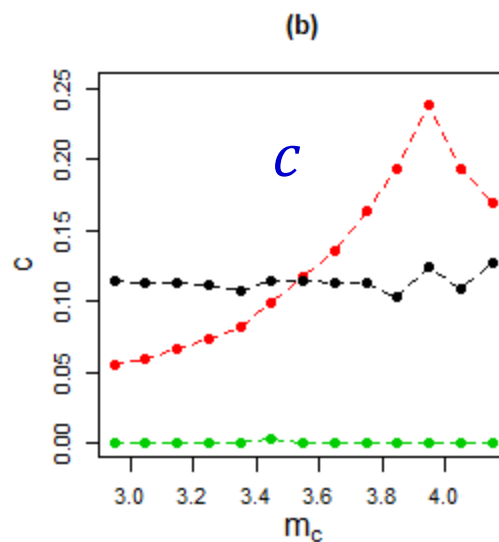
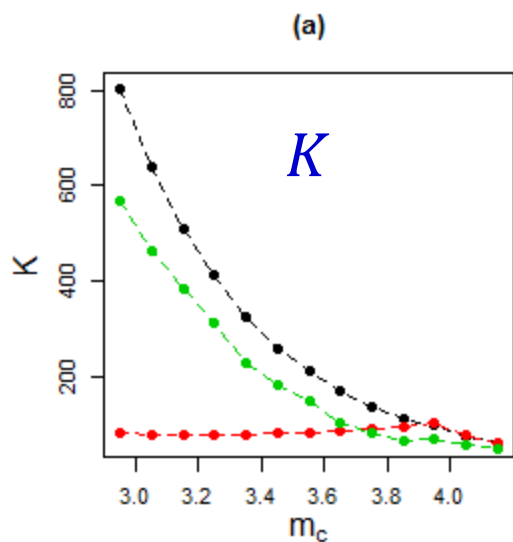
Cumulative freq. vs time



# Influence of short-term aftershock missing on estimating the Omori formula

Omori-Utsu formula:

$$\lambda(t) = \frac{K(p-1)}{c} \left(1 + \frac{t}{c}\right)^{-p}, \quad t > 0$$



●●●●●●●● Replenished data

●●●●●●●● Observed data  
Start fitting at 0

●●●●●●●● Observed data  
Start fitting at 54<sup>th</sup> day

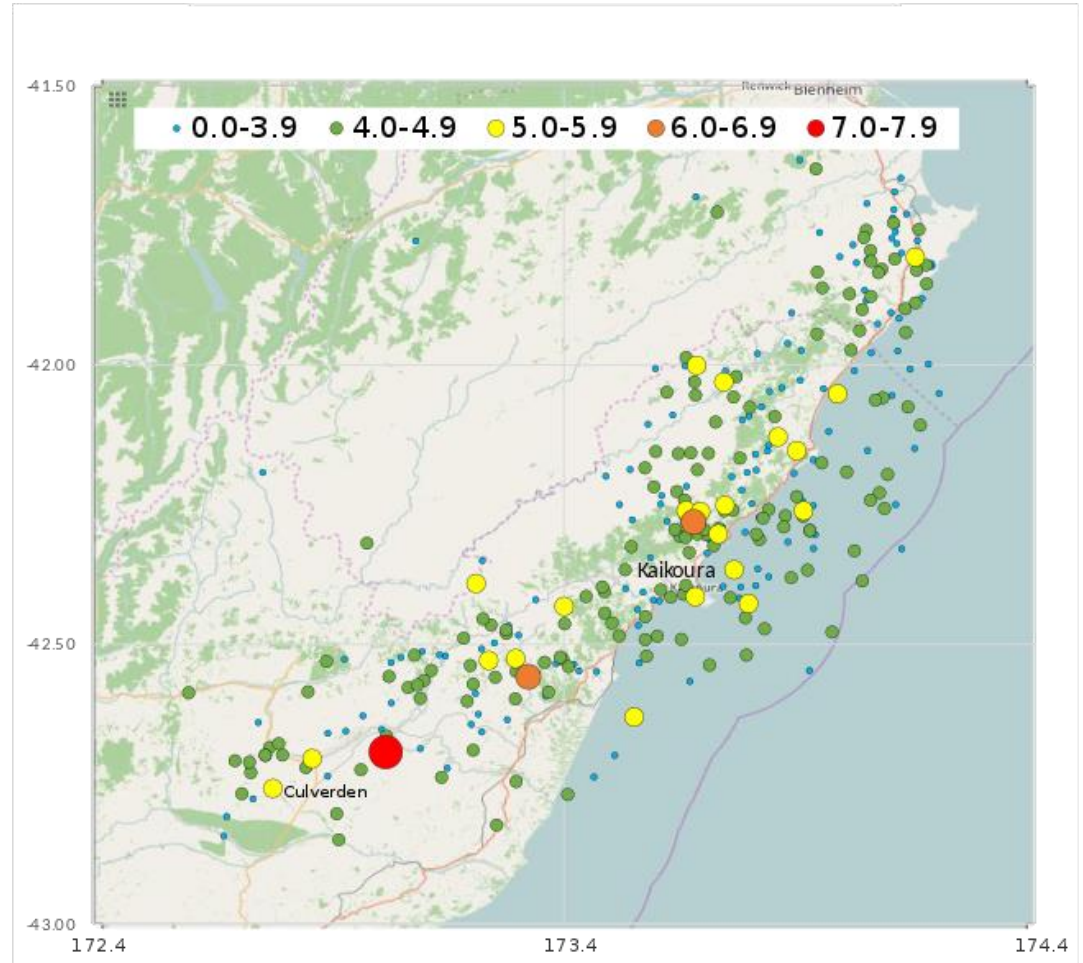
# Application 3: the Kaikoura aftershock sequence data

## Kaikoura EQ: Mw7.8 2016/11/14

Data selection:

Time: 2014/1/1~2017/2/17

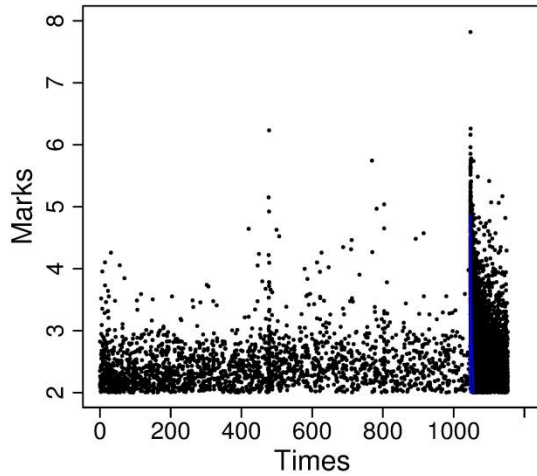
Magnitude: 2.0+



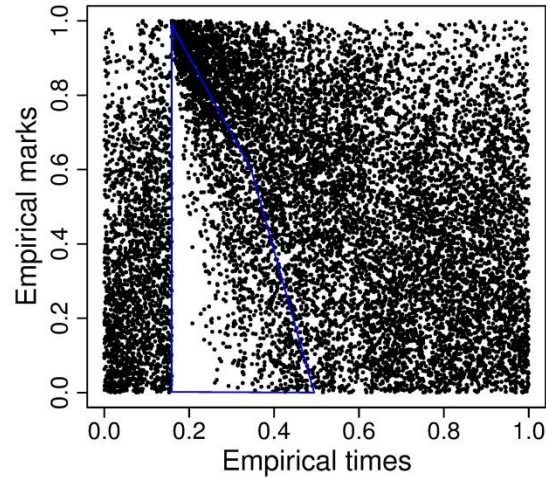


# Application 3: the Kaikoura aftershock sequence data

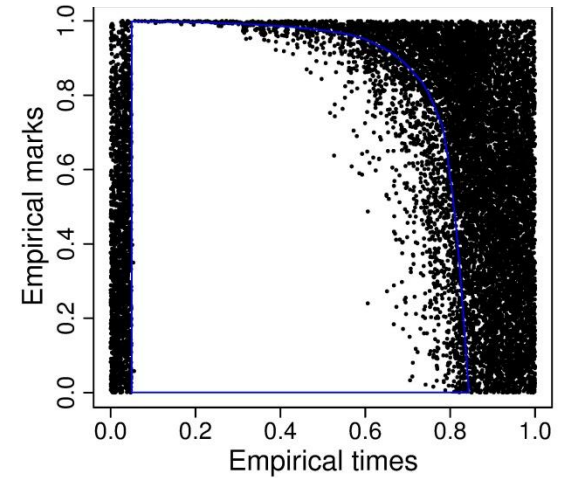
Observed data



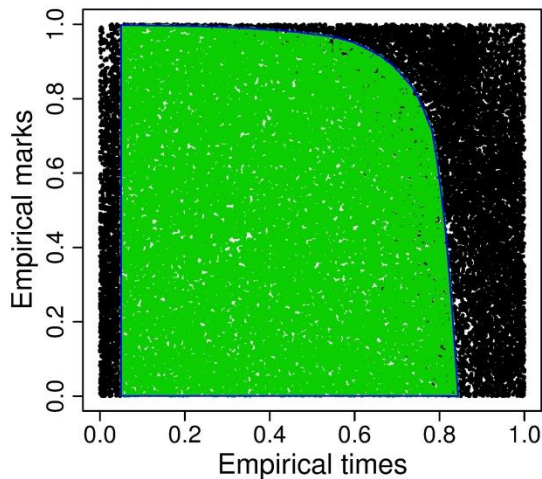
Biscale transformed



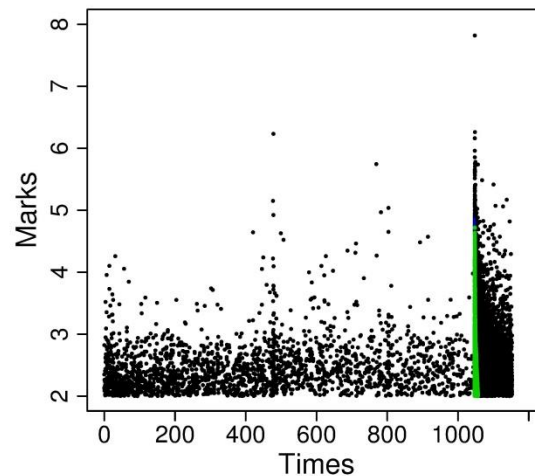
Estimate transformation under complete data



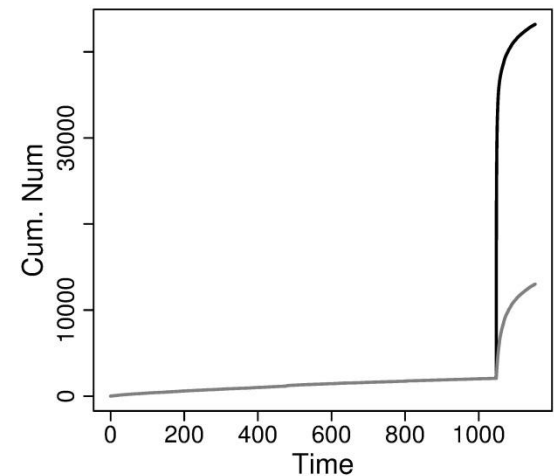
Transformed replenished data




Original replenished data



Cumulative freq. vs time





# **On the biased estimate of earthquake clustering parameters caused by short-term aftershock missing**

**Errors in estimated parameters propagate in forecasting.**

# The ETAS model

- **Conditional intensity** (Ogata, 1998)

$$\begin{aligned}\lambda(t, x, y) &= E[N(dtdxdy) | H_t] / |dtdxdy| \\ &= \nu\mu(x, y) + \sum_{i:t_i < t} \kappa(m_i) g(t - t_i) f(x - x_i, y - y_i, m_i)\end{aligned}$$

1. Direct productivity:  $\kappa(m) = Ae^{\alpha(m-m_c)}, \quad m \geq m_c$

2. Time p.d.f (Omori-Utsu):  $g(t) = (p-1)(1+t/c)^{-p} / c, \quad t > 0$

3. Location p.d.f:  $f(x, y | m) = \frac{q-1}{\pi D e^{\gamma(m-m_c)}} \left( 1 + \frac{x^2 + y^2}{D e^{\gamma(m-m_c)}} \right)^{-q}$

**Temporal version** (Ogata 1988, JASA)

$$\lambda(t) = \mu + K \sum_{i:t_i < t} \frac{e^{\alpha(m_i - m_c)}}{(t - t_i + c)^p}$$

$$A = \frac{Kc^{1-p}}{p-1} = \int_0^\infty \frac{K}{(t+c)^p} dt$$

expected # of direct offspring from  $m_c$



# What influence the estimate of the ETAS parameters



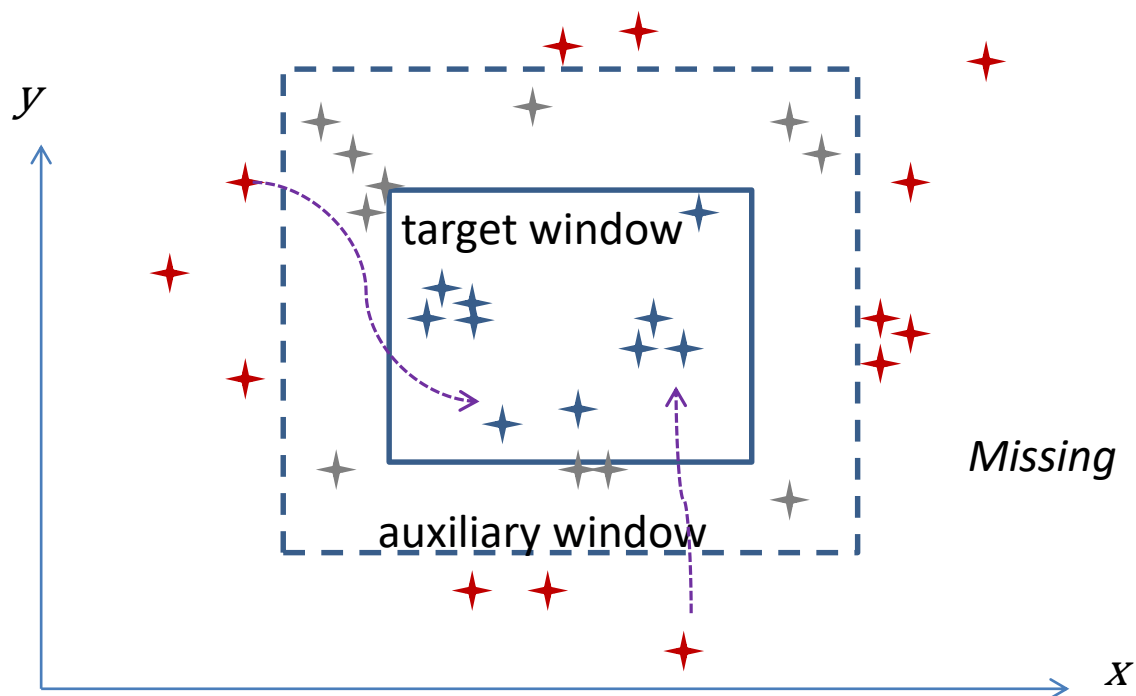


# What influence the estimate of the ETAS parameters

*Missing links!*

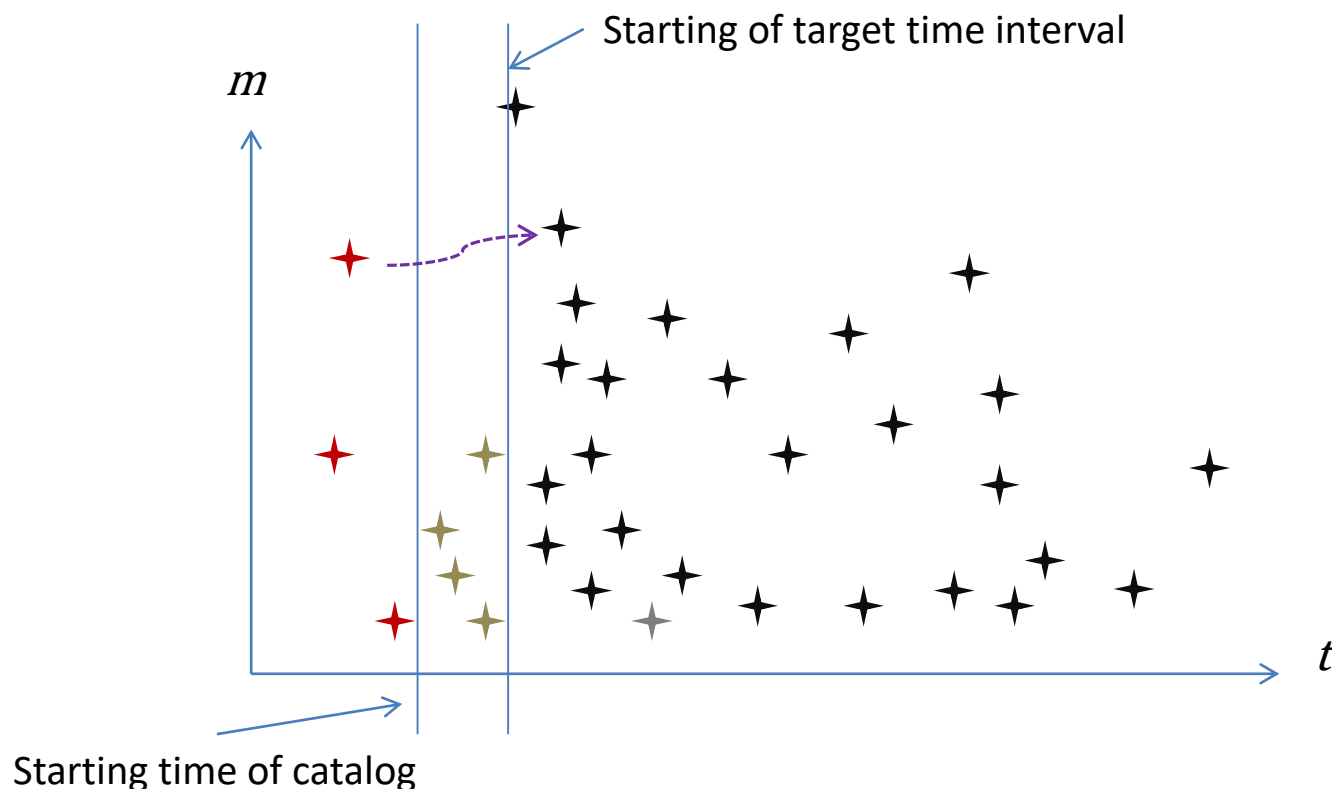
# What influence the estimate of the ETAS parameters: *Missing links!*

## 1. Missing links in space



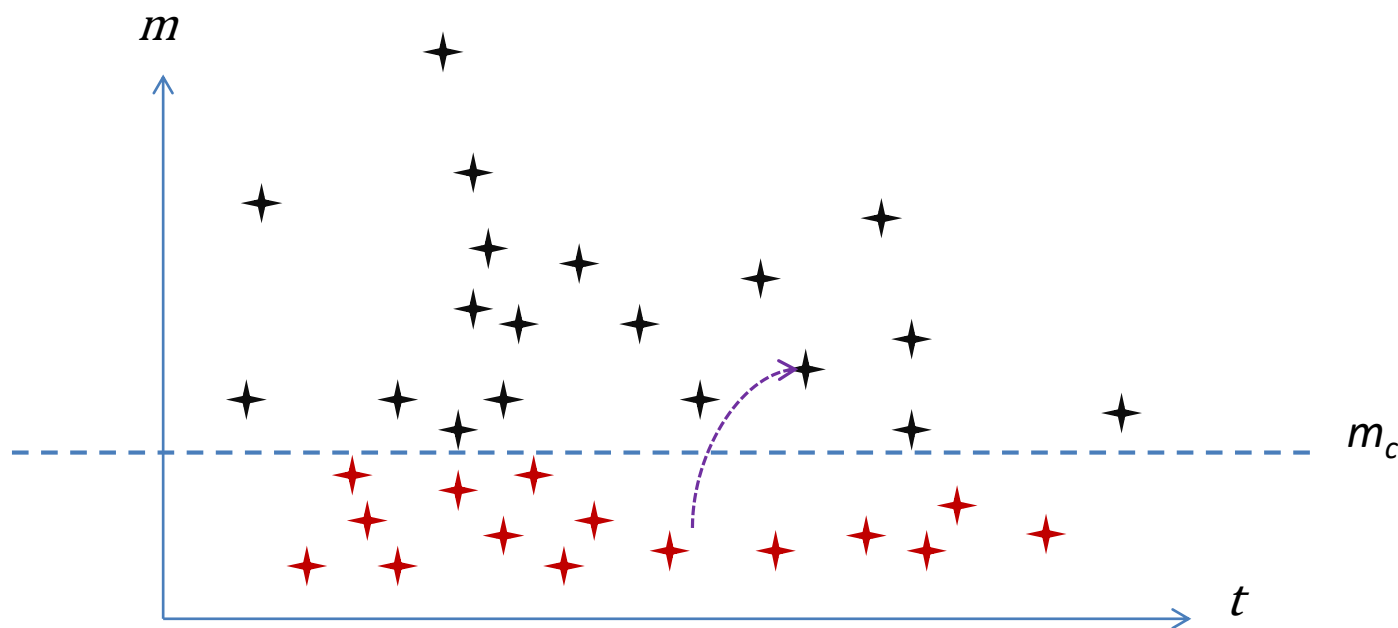
# What influence the estimate of the ETAS parameters: *Missing links!*

## 2. Missing links in time



# What influence the estimate of the ETAS parameters: *Missing links!*

## 3. *Missing links in magnitude*





# The influence of the above missing links has been studied in

- Sornette & Werner (2005) **JGR**, 110, B09303.
- Sornette & Werner (2005) **JGR**, 110, B08304.
- Wang, et al. (2010) **BSSA**, 100 , 1989 – 2001.
- Wang, Jackson & Zhuang. (2013) **GRL.**, 37 , L21307.
- etc..



# What influence the estimate of the ETAS parameters

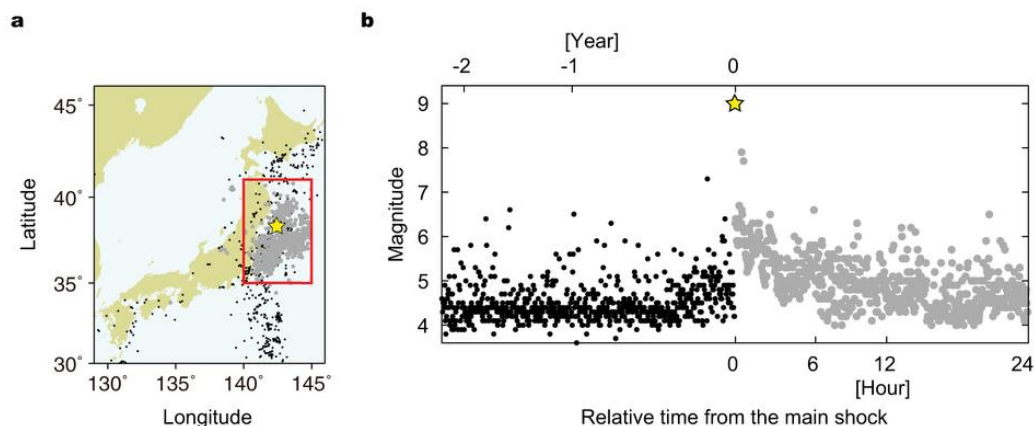
*4. Missing links caused by short-term missing of aftershocks.*



# What influence the estimate of the ETAS parameters

## *4. Missing links caused by short-term missing of aftershocks.*

Missing events after the mainshock (from Omi etc, 2013)



# Previous studies for fixing the problems of short-term missing aftershocks

## Observational approaches

- ◆ waveform-based earthquake detection methods (e.g., Enescu et al., 2007, 2009; Peng et al., 2007; Marsan and Enescu, 2012; Hainzl, 2016).
- ◆ Energy based description (Sawazaki and Enescu, 2014)

## Statistical approaches

- ◆ (Ogata, Omi, Iwata) Bayesian, assuming GR relation for whole range
- ◆ (Marsan and Enescu, 2012) Assuming Omori-Utsu formula or ETAS model
- ◆ This study: Conditional independence between magnitudes and occurrence times

# Application to the Kumamoto aftershock sequence data

## Data Selection

Time: 2016/4/1~2016/4/21

Mag.: 1.0+

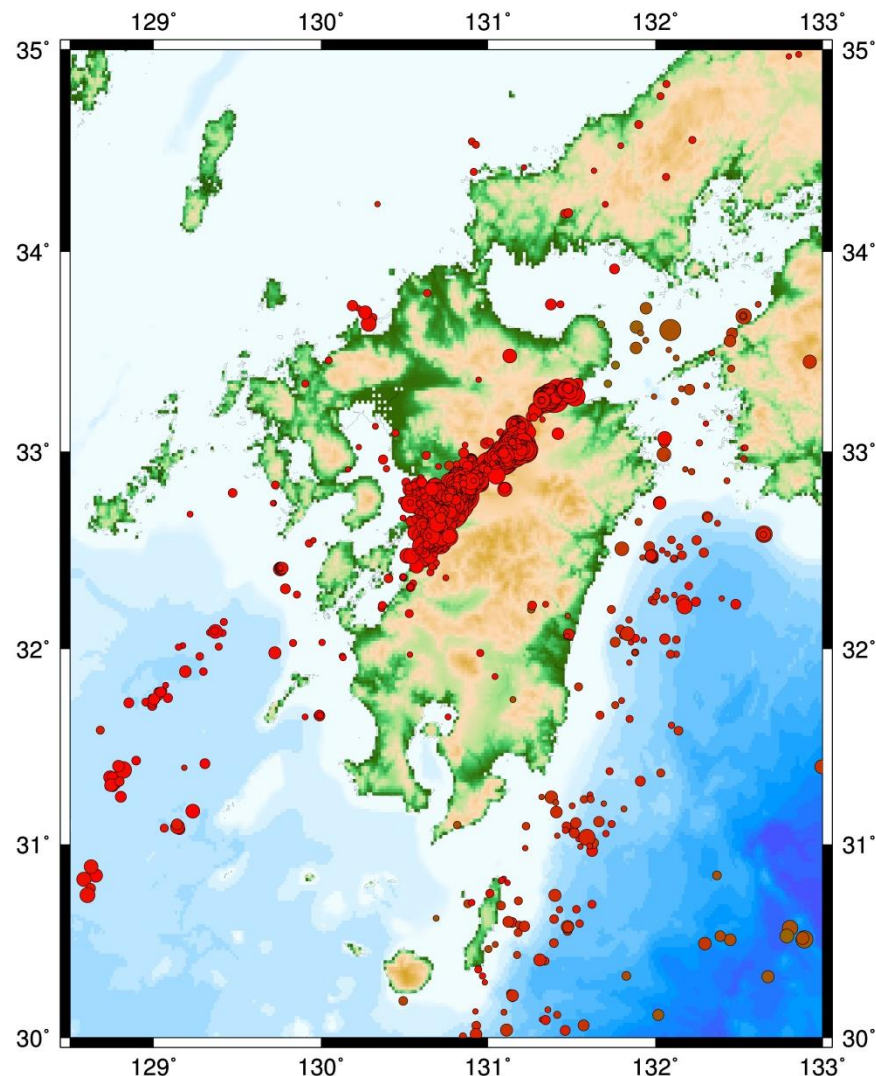
Depth: < 100 km

Space: 128° -- 133°  
30° -- 35°

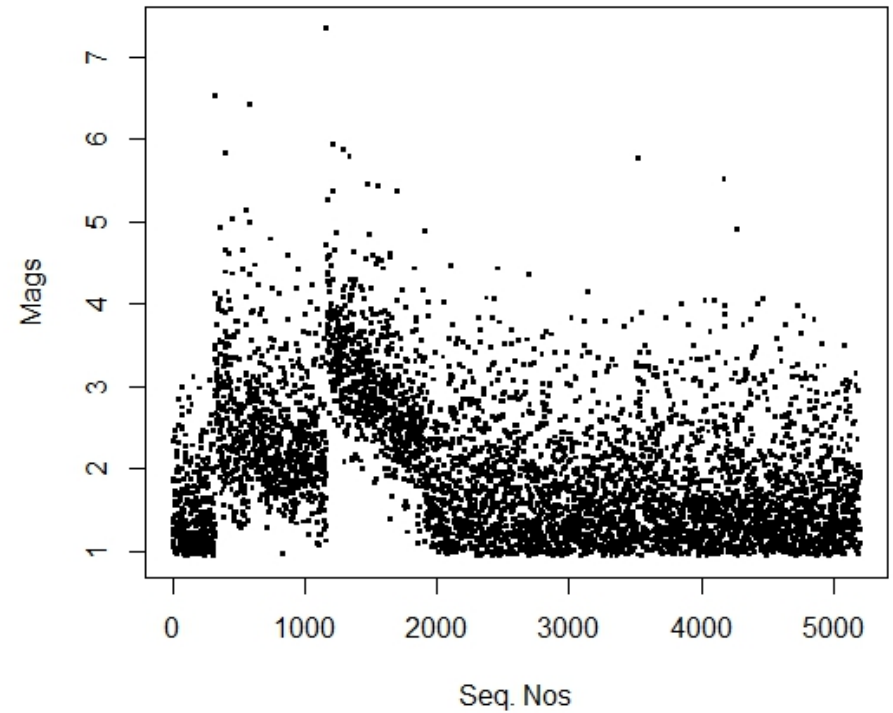
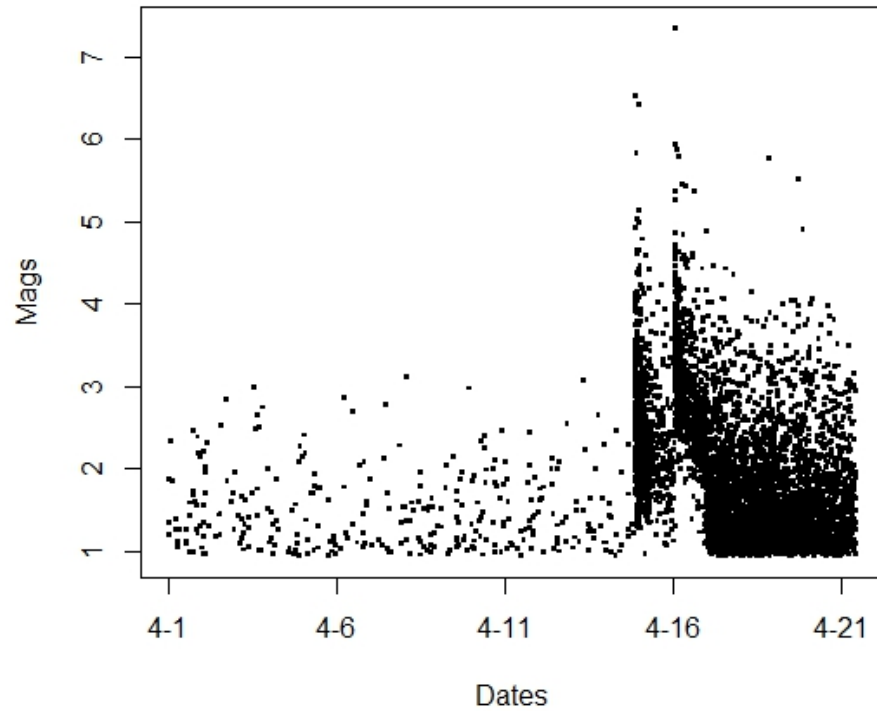
**2016-04-14 21:26 (130.81 32.74) M6.5**

**2016-04-15 00:03 (130.78 32.70) M6.4**

**2016-04-16 01:25 (130.76 32.75) M7.3**



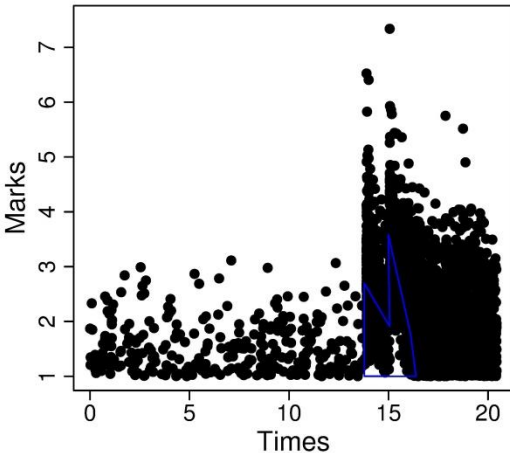
# Application to the recent Kumamoto aftershock sequence data



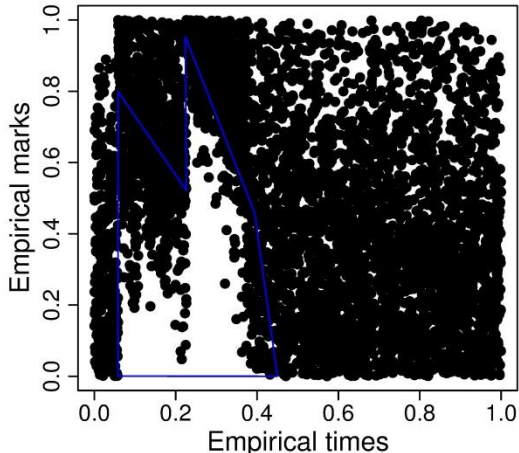
# Replenish the missing data

Method (Zhuang et al, 2016)

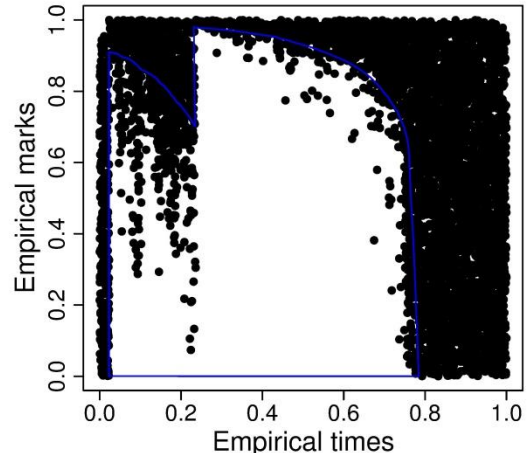
Observed data



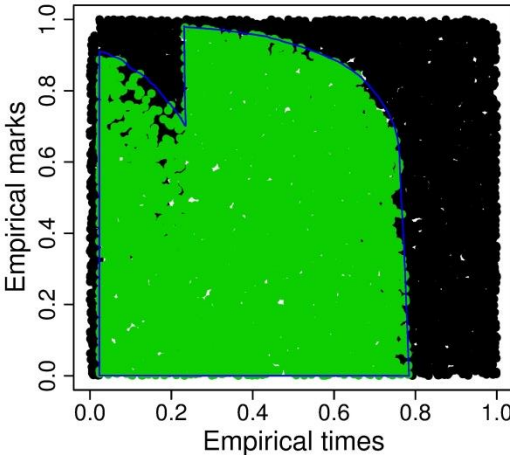
Biscale transformed



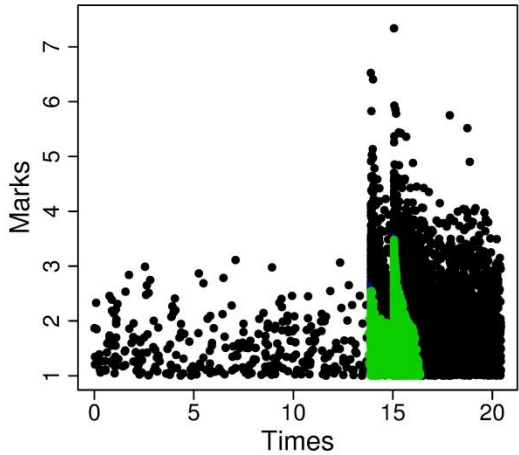
Estimated transformation under complete data



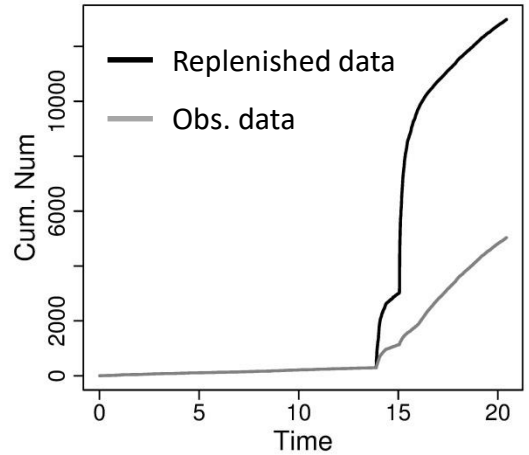
Transformed replenished data



Original replenished data



Cumulative freq. vs time



## Consider only the temporal ETAS model

$$\lambda(t) = \mu + K \sum_{i:t_i < t} \frac{e^{\alpha(m_i - m_c)}}{(t - t_i + c)^p} \quad (\text{Ogata 1988, JASA})$$

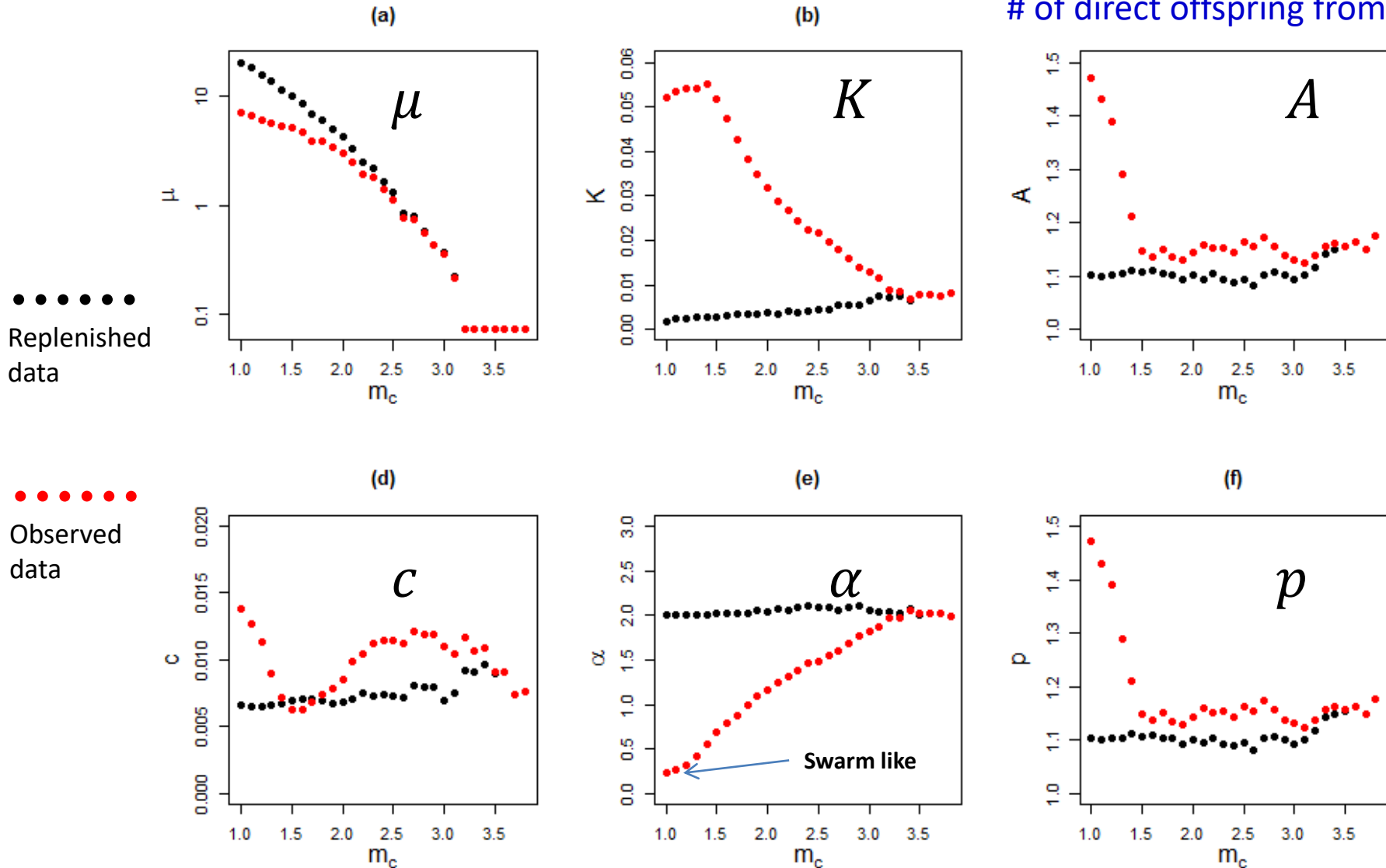
$$A = \frac{Kc^{1-p}}{p-1} = \int_0^\infty \frac{K}{(t+c)^p} dt \quad : \quad \text{expected \# of direct offspring from } m_c$$



# Influence of short-term aftershock missing on estimating the ETAS model

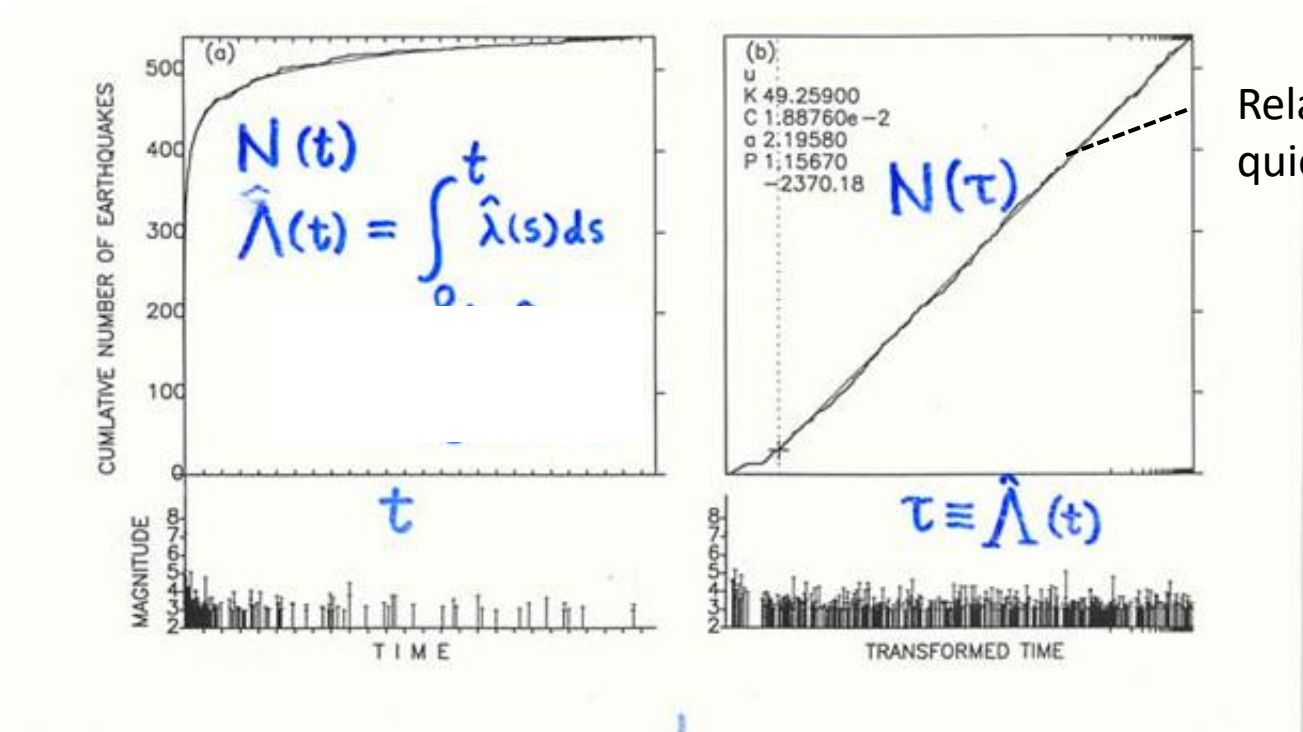
$$A = \frac{Kc^{1-p}}{p-1}$$

# of direct offspring from  $m_c$



# Quiescence related to the ETAS model

## . Transformed Time sequence (Ogata, 1992, JGR)

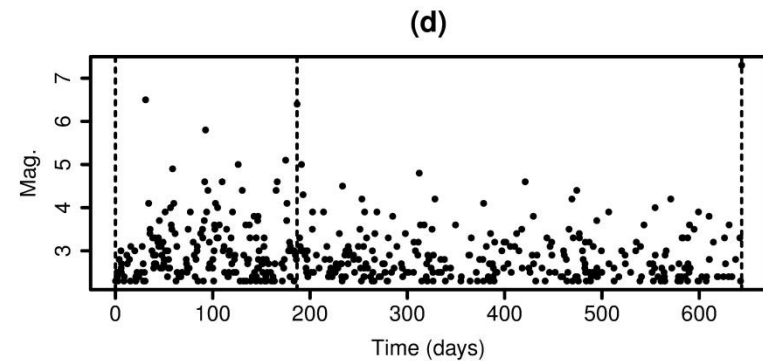
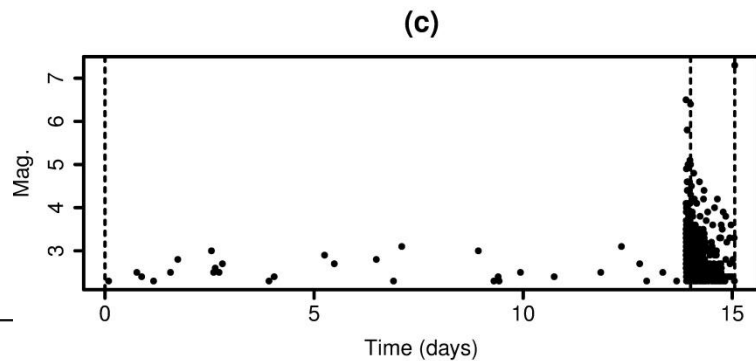
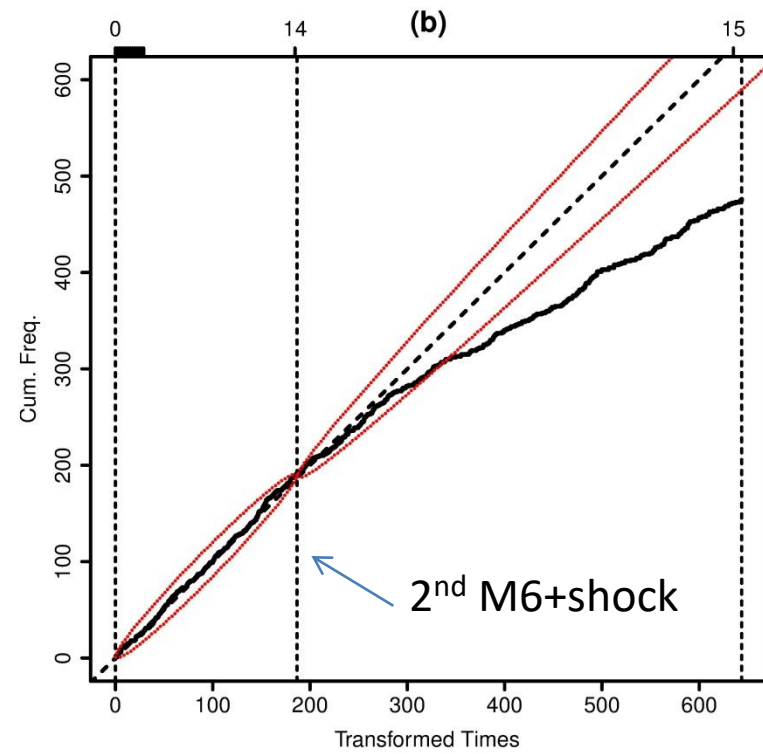
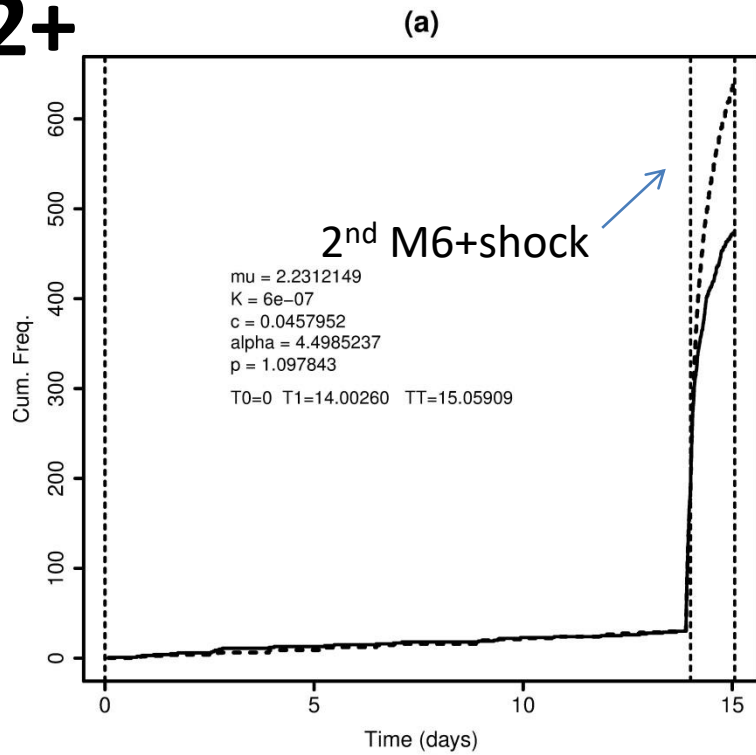


$$t_i \rightarrow \tau_i = \int_0^{t_i} \lambda(u) du$$

If  $\{t_i\}$  is the observation of a process determined by conditional intensity  $\lambda(t)$ , the  $\{\tau_i\}$  is a standard Poisson process.

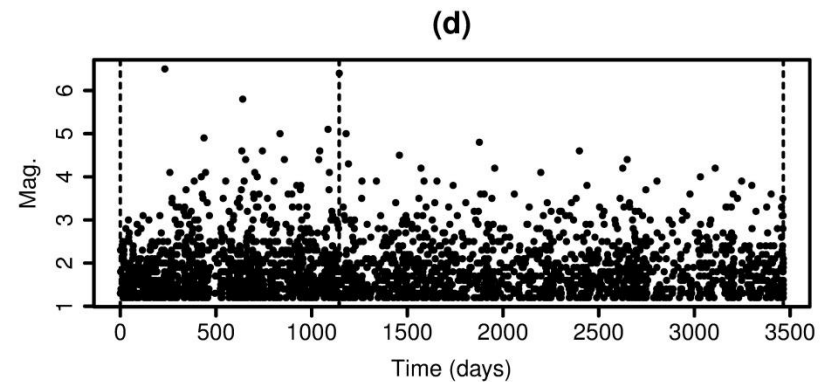
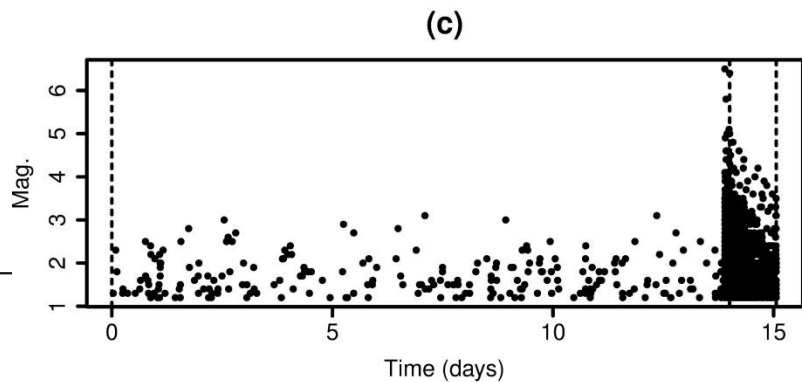
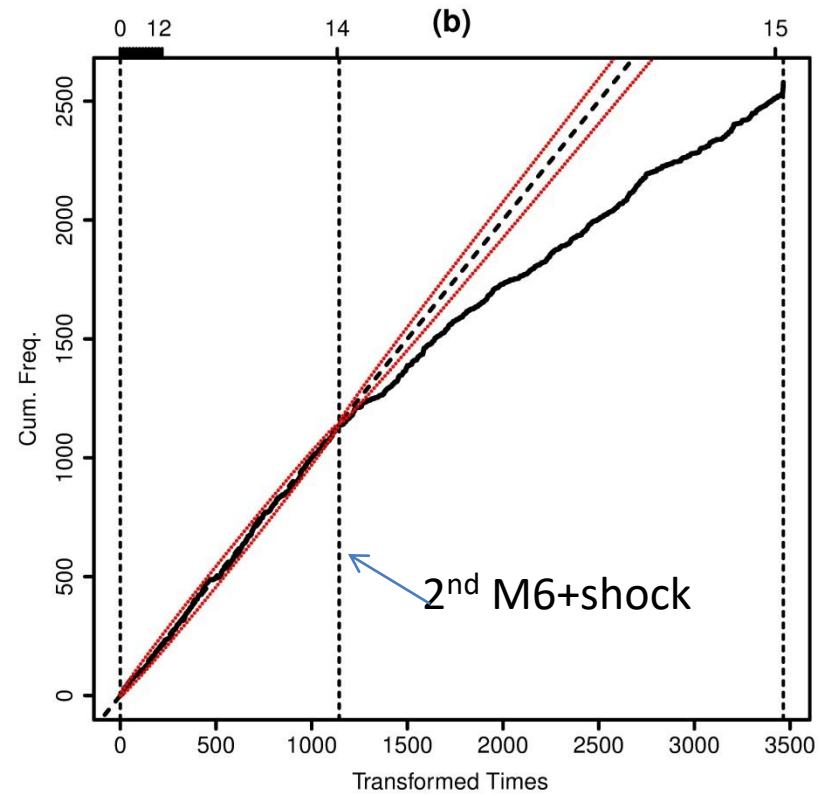
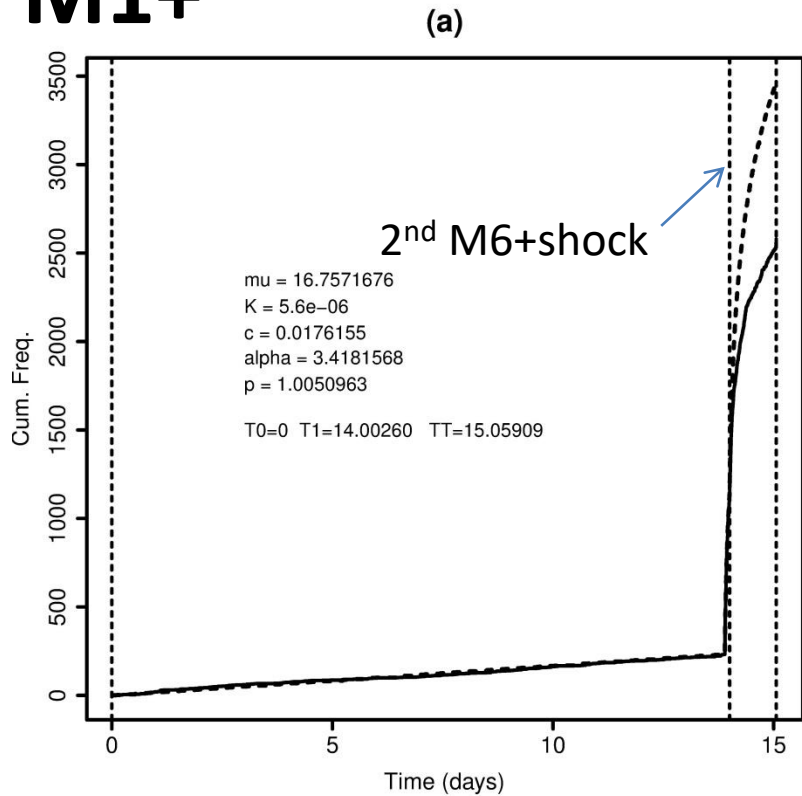
# Relative quiescence --- original dataset

## M2+



# Relative quiescence --- Replenished dataset

## M1+



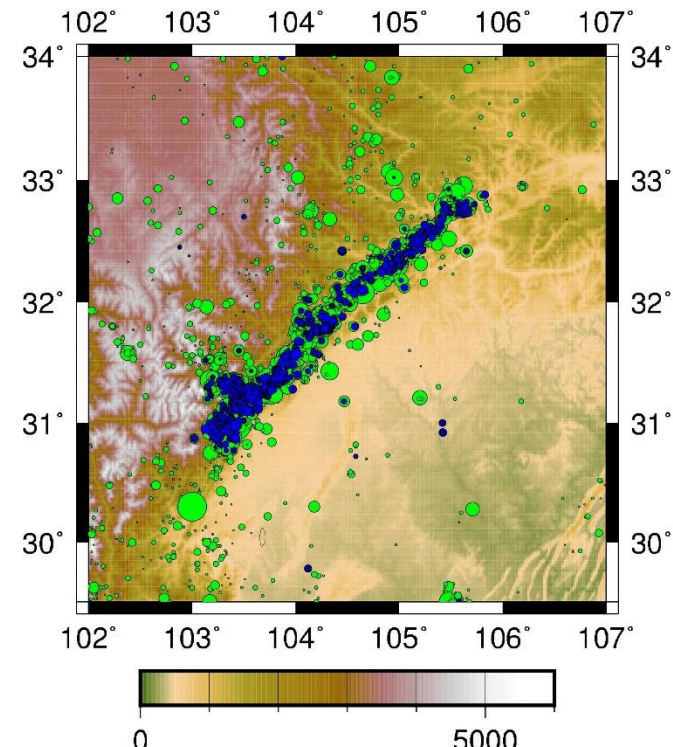
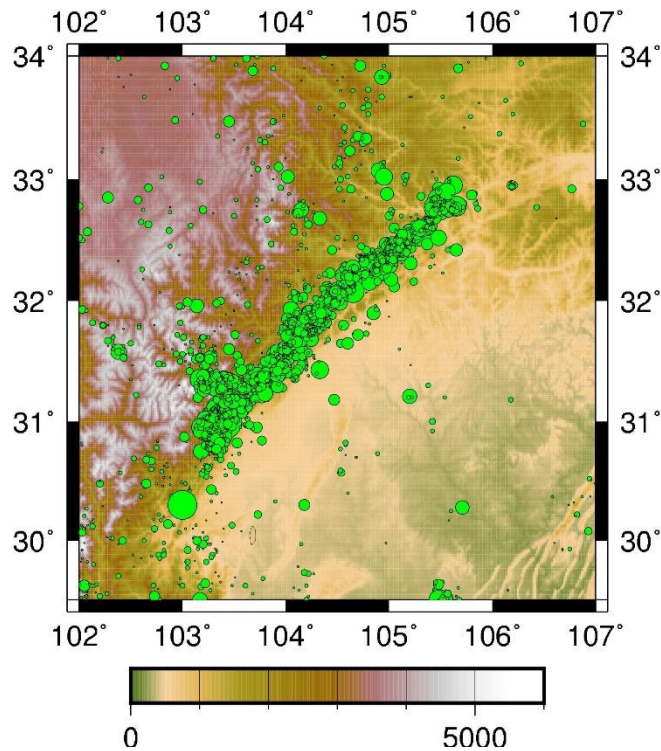
# Conclusions

- A method for replenishing missing data in marked temporal point process, which only makes use of the assumption that the marks and occurrence times of events are independent or that the mark distributions has no dramatic changes along the time axis, regardless of how the events interact on the time axis.
- The key point is an iterative algorithm for estimating the missing area in the transform domain according to the parts where data are completely recorded.
- This method is applied to the eruption of the Hakone volcano in Japan and the earthquake catalogue from Southwest China including the aftershock zone of the 2008 M7.8 Wenchuan earthquake



# Future Researches

- Extending to higher dimension or spatiotemporal cases.
- If available, more information of the marks or times can be used to improve the replenishment. For example, distribution of earthquake magnitudes is well known (G-R law).







*Thank you for  
listening.*